

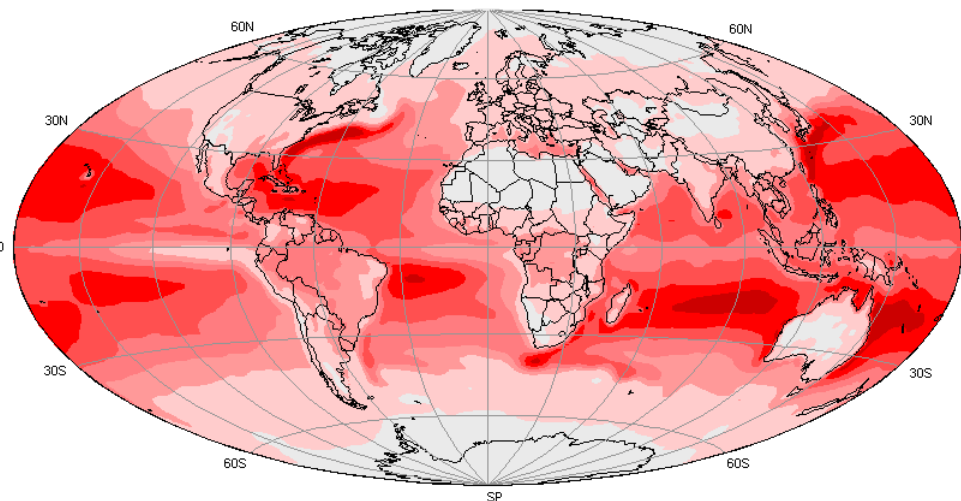
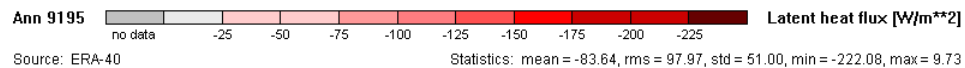
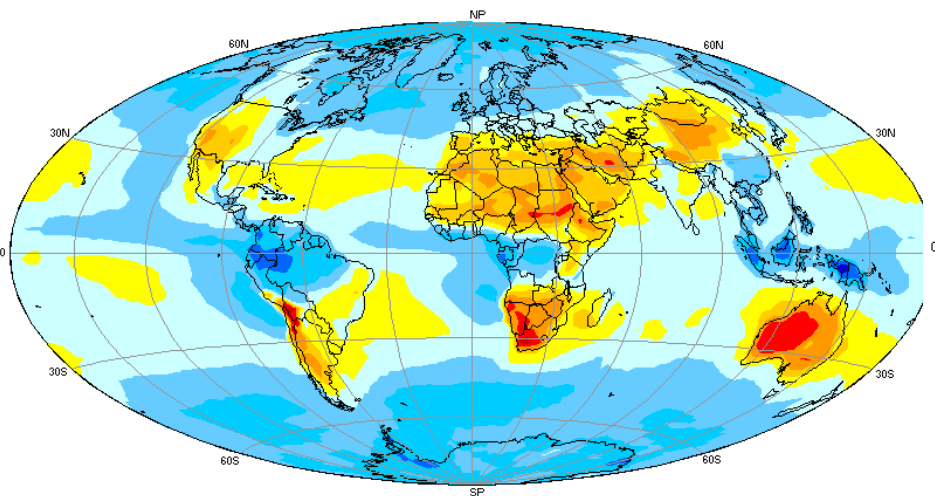
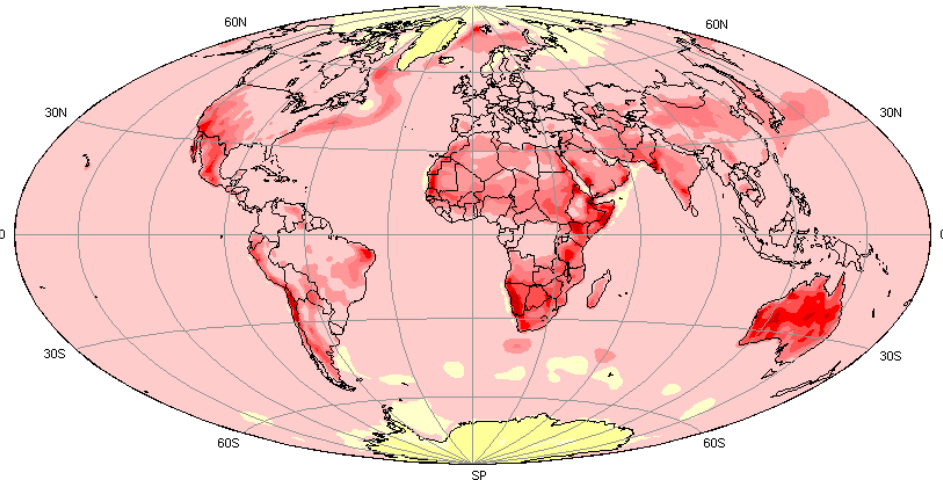
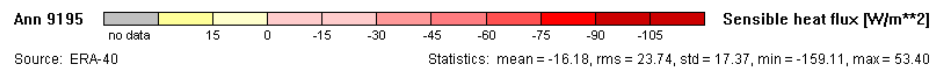
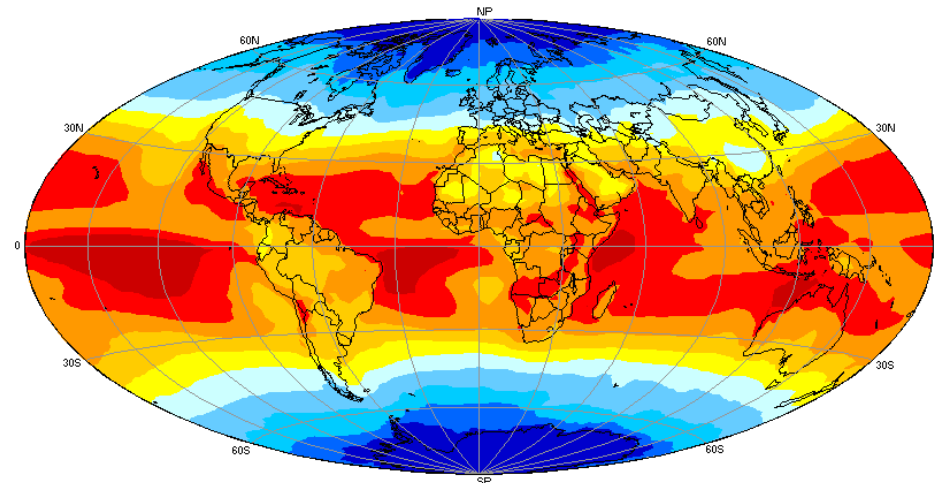
Implicit coupling between the atmospheric boundary layer and the continental surface for solving the surface temperature

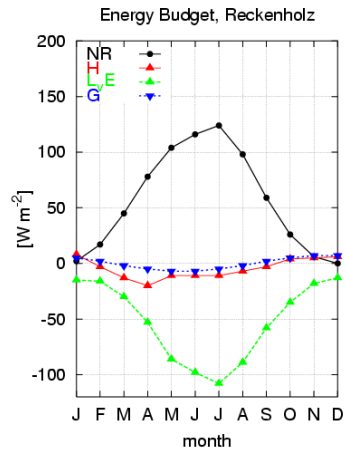
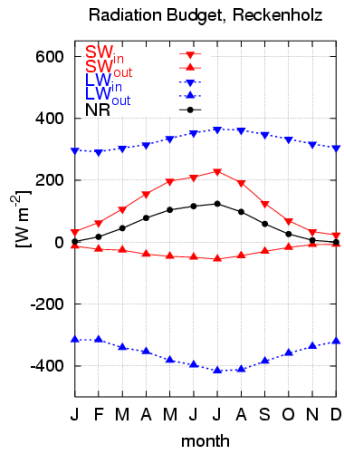
F Cheruy ,  
F. Wang , J.L Dufresne , J. Polcher , A. Ducharne

ORCHIDEE Formation, 7 Novembre 2014

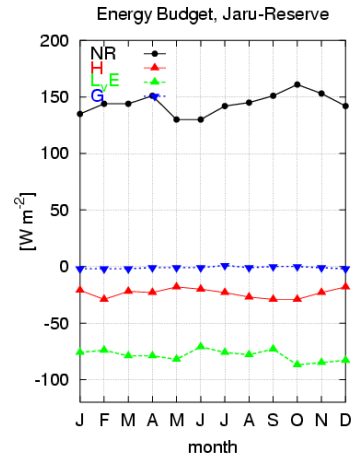
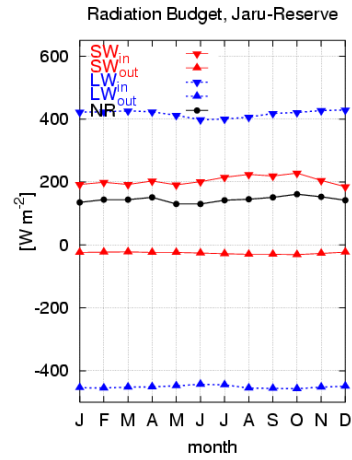
- Introduction
- Surface energy budget equation
- Turbulent fluxes
- Radiation
- Solving for the Surface temperature
- Routines used

$$SW_{net} + LW_{net} + H + L.E + G = 0$$



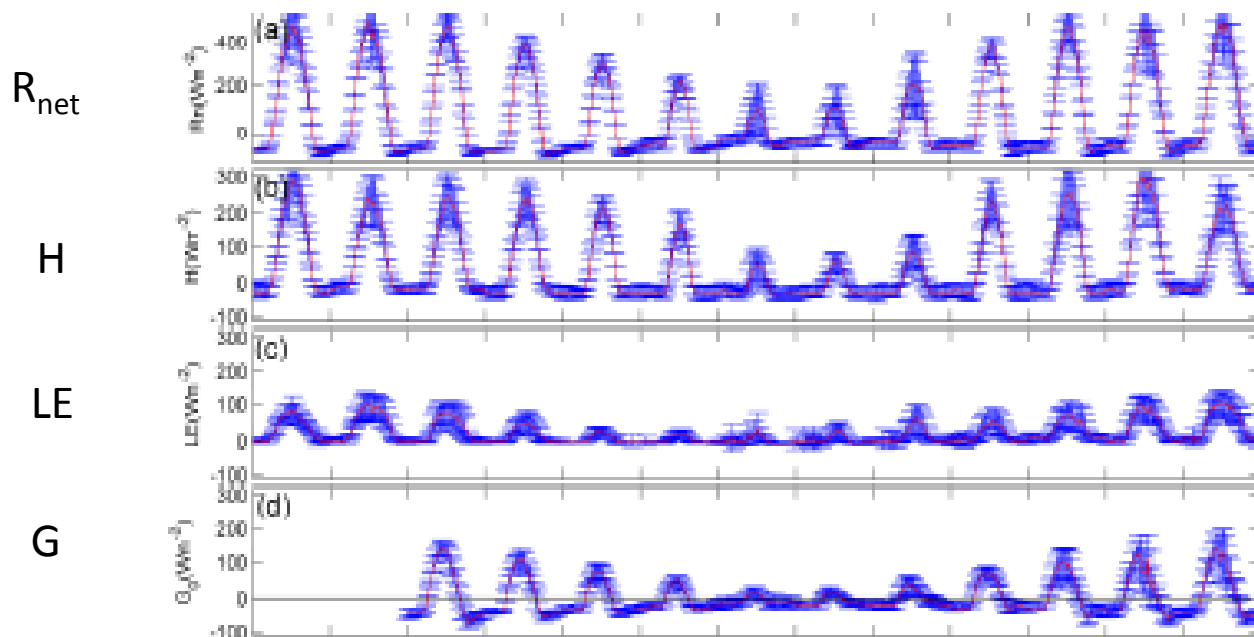
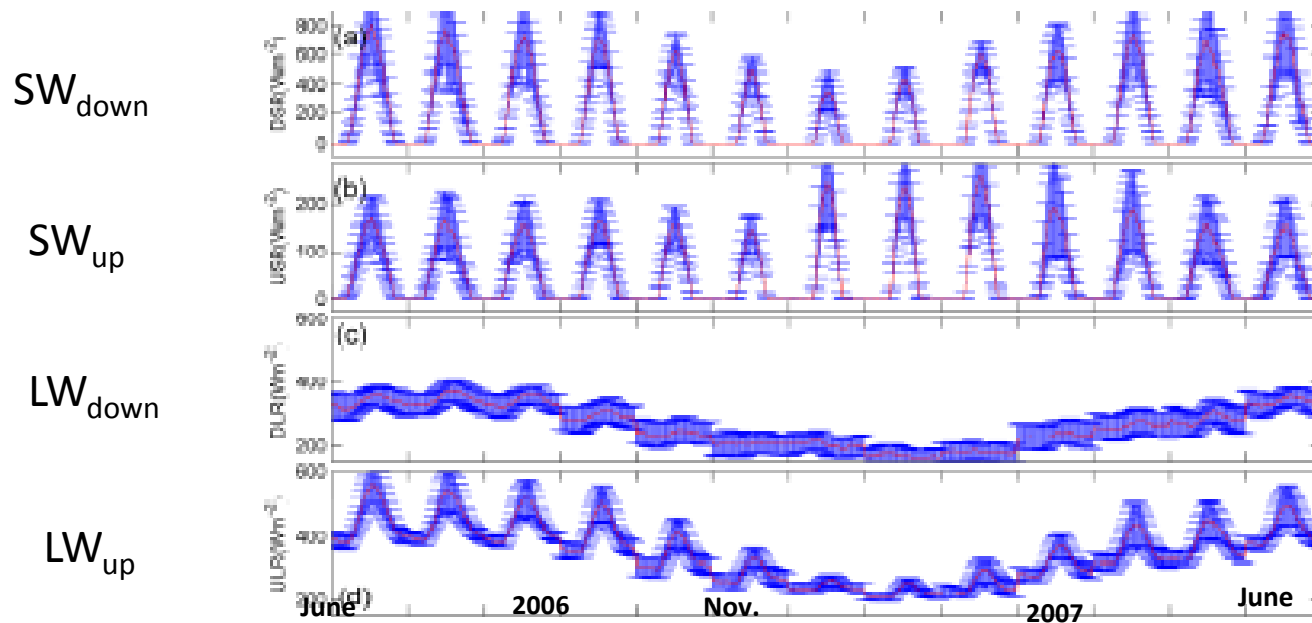


Grassland in Switzerland



Tropical Rain Forest

# Diurnal variations of weekly mean of surface energy budget terms (44N, 116E, Steppe, China)



## Surface energy budget

$$C_s \frac{\partial}{\partial t} T_s = SW_{net} + LW_{net} + F_{H,s}^t + L.E + G \quad (1)$$

$$LW_{net} = LW_{down} - \varepsilon \sigma T_s^4 - (1 - \varepsilon) LW_{down} \quad (2)$$

$$F_{H,s}^t = \rho |\vec{v}| C_d (H_1^t - H_s^t) \quad (3)$$

$$L.E_s^t = L\rho |\vec{v}| C_d \beta (q_1^t - q_{sat}(T_s^t)) \quad (4)$$

Implicit coupling: The atmospheric temperature, humidity and surface conditions are estimated synchronously (at the same time step)

$$C_s \frac{(T_s^t - T_s^{t-\delta t})}{\delta t} = SW_{net}^t + LW_{net}^t + H^t + L.E^t + G^t \quad (5)$$

$$C_s \frac{(T_s^t - T_s^{t-\delta t})}{\delta t} = SW_{net} + LW_{down} + G^t - \epsilon\sigma * T_s^{t4} + H^t + L.E^t \quad (6)$$

$G^t$ , computed by thermosoil

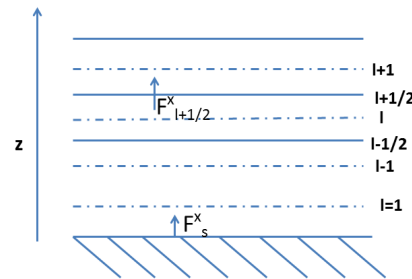
$C_s$  is the apparent surface heat capacity and is computed by thermosoil

## Turbulent fluxes and their dependence on the surface temperature

The diffusion equation of the potential enthalpie or the specific humidity (X) in the boundary layer  $F^X = -\rho k \frac{\partial}{\partial z} X$  and  $\frac{\partial}{\partial t} X = -\frac{1}{\rho} \frac{\partial}{\partial z} F^X$  and the continuity equation are discretized in space and time.

$$\frac{X_l^t - X_l^{t-\delta t}}{\delta t} = \frac{-g}{\delta P_l} (F_{l+1/2}^X - F_{l-1/2}^X) \quad (7)$$

$$F_{l-1/2}^X = -\rho_{l-1/2} k_{l-1/2} \frac{(X_l^t - X_{l-1}^t)}{z_l - z_{l-1}} \quad (8)$$



$$F_{X,s}^t = \rho |\vec{v}| C_d (X_1^t - X_s^t) \quad (9)$$



For the interface variable this can be re-written as

$$X_1^t = A_X^1 + B_X^1 \cdot F_{s,X}^t \delta t \quad (10)$$

$$X_2^t = C_2^X + D_2^X H_1^t \quad (11)$$

$$F_{X,s}^t = \rho |\vec{v}| C_d (X_1^t - X_s^t) \quad (12)$$

$$A_X^1 = \frac{X_1^{t-\delta t} \delta P_1 + S_2^X C_2^X}{\delta P_1 + S_2^X (1 - D_2^X)} \quad (13)$$

$$B_X^1 = \frac{-g}{\delta P_1 + S_2^X (1 - D_2^X)} \quad (14)$$

with  $S_2 = \frac{g^2 \rho_{3/2}^2 k_{3/2} \delta t}{P_1 - P_2}$  and  $A_X^1$  and  $B_X^1$  do not contain any information on the surface condition at the time  $t$  and can be computed from the PBL prior to the land surface processes.

## Sensible heat flux

$$H_1^t = A_H^1 + B_H^1 \cdot F_{s,H}^t \delta t \quad (15)$$

$$F_{s,H}^t = \frac{1}{zik t} (H_1^t - H_s^t) \quad (16)$$

with  $H = C_p T \left(\frac{P_r}{P}\right)^\kappa$ ,  $H_s = C_p T_s$  and  $\frac{1}{zik t} = \rho |\vec{v}| C_d$

$$F_{s,H}^t = \frac{1}{zik t} (A_H^1 + B_H^1 \cdot F_{s,H}^t \delta t - H_s^t) \quad (17)$$

$$F_{s,H}^t = \frac{1}{zik t} \left[ \frac{(A_H^1 - H_s^{t-\delta t})}{1 - \frac{1}{zik t} B_H^1 \delta t} - \frac{(H_s^t - H_s^{t-\delta t})}{1 - \frac{1}{zik t} B_H^1 \delta t} \right] \quad (18)$$

$$F_{s,H}^t = sensfl_{old} - sensfl_{sns} (T_s^t - T_s^{t-\delta t}) \quad (19)$$

➤  $sensfl_{old}$  is based on the surface condition at the previous time step ( $t - \delta t$ ) and the atmospheric condition at the time step  $t$

➤  $sensfl_{sns}$  is a sensitivity term to the variation of surface temperature. ➤ In the ORCHIDEE code,  $A_H^1$  is called `petBcoef`,  $B_H^1 \delta t$  is called `petAcoef`

## Latent heat flux (evaporation)

$$F_{s,q}^t = L\rho |\vec{v}| C_d \beta (q_1^t - q_{sat}(T_s^t)) \quad (20)$$

$$q_{sat}(T_s^t) = q_{sat}(T_s^{t-\delta t}) + \frac{\partial q_{sat}}{\partial t} \Big|_{(T=T_s^{t-\delta t})} (T_s^t - T_s^{t-\delta t}) \quad (21)$$

$$q_1^t = A_q^1 + B_q^1 \cdot F_{s,q}^t \delta t \quad (22)$$

$$F_{s,q}^t = \frac{1}{zik t} \beta (A_q^1 + B_q^1 \cdot F_{s,q}^t \delta t - q_{sat}(T_s^{t-\delta t})) + \frac{\partial q_{sat}}{\partial t} \Big|_{(T=T_s^{t-\delta t})} (T_s^t - T_s^{t-\delta t}) \quad (23)$$

$$F_{s,q}^t = \frac{\frac{1}{zik t} \beta (A_q^1 - q_{sat}(T_s^{t-\delta t}))}{(1 - \frac{1}{zik t} \beta B_q^1 \delta t)} - \frac{\frac{1}{zik t} \beta \frac{\partial q_{sat}}{\partial t} \Big|_{(T=T_s^{t-\delta t})}}{(1 - \frac{1}{zik t} \beta B_q^1 \delta t)} (T_s^t - T_s^{t-\delta t}) \quad (24)$$

In the ORCHIDEE code,  $A_q^1$  is called `peqBcoef`,  $B_q^1 \delta t$  is called `peqAcoef`

$$L \cdot F_{s,q}^t = L.lareva_{old} - \frac{L}{C_p} lareva_{sns} (H_s^t - H_s^{t-\delta t}) \quad (25)$$

## Radiation

$$netrad = SW_{net} + LW_{down} - \varepsilon\sigma * T_s^{t4} - (1 - \varepsilon)LW_{down} \quad (26)$$

In the present version the LW emissivity is set one

$$netrad = SW_{net} + LW_{down} - (1 - \varepsilon)LW_{down} - \varepsilon\sigma * T_s^{t-\delta t4} - 4\varepsilon\sigma T_s^{t-\delta t3} (T_s^t - T_s^{t-\delta t}) \quad (27)$$

$$netrad = netrad_{old} + netrad_{sns}(T_s^t - T_s^{t-\delta t}) \quad (28)$$

surface temperature at the new time step

$$sum_{old} = netrad_{old} + sensfl_{old} + lareva_{old} \quad (29)$$

$$sum_{sns} = netrad_{sns} + sensfl_{sns} + lareva_{sns} \quad (30)$$

$$C_s \frac{(T_s^t - T_s^{t-\delta t})}{\delta t} = sum_{old} + G + sum_{sns}(T_s^t - T_s^{t-\delta t}) \quad (31)$$

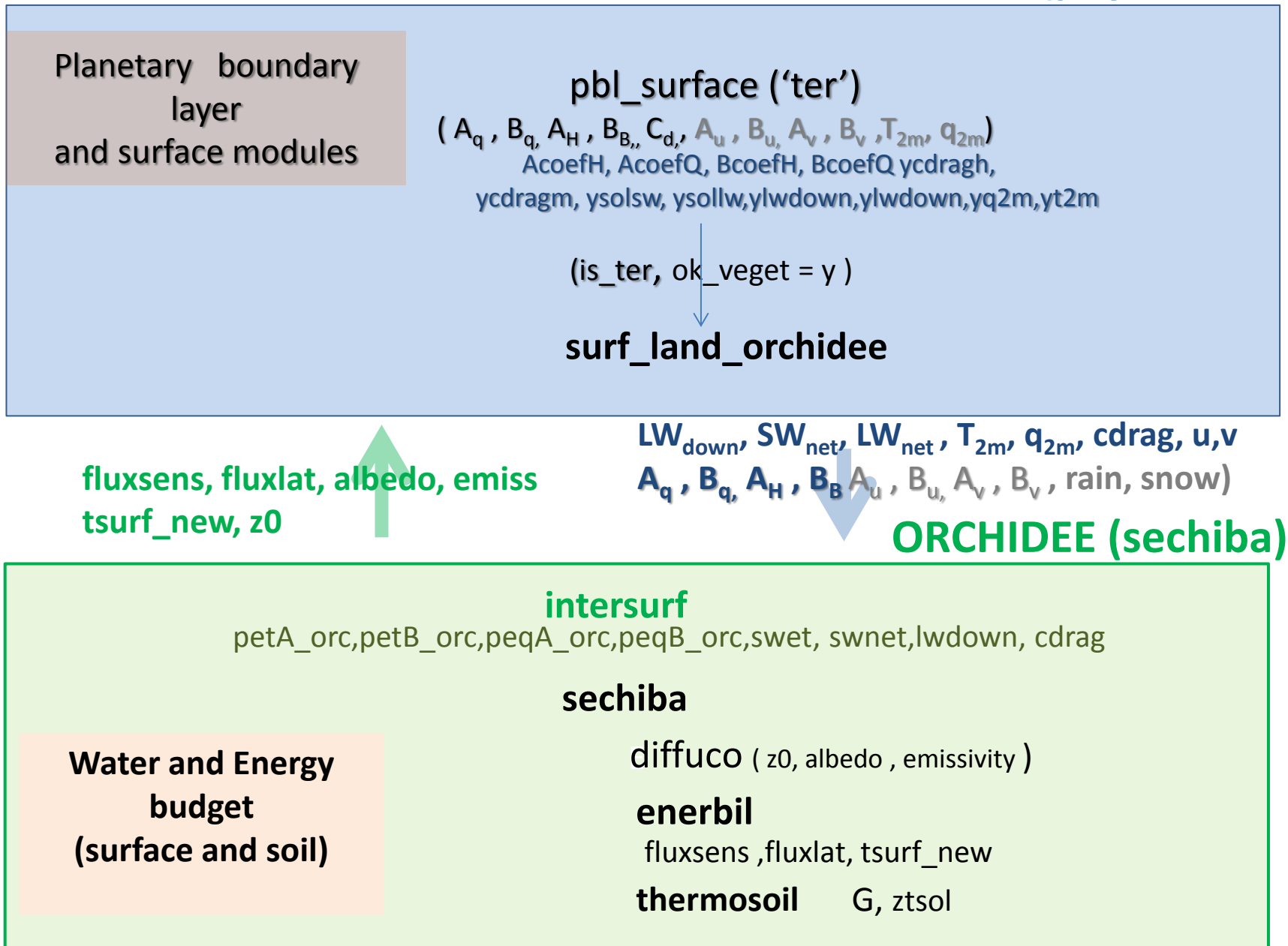
$$(T_s^t - T_s^{t-\delta t}) = \frac{\delta t(sum_{old} + G)}{C_s - sum_{sns}\delta t} \quad (32)$$

About the radiation the time step of the radiation is usually longer than the one of ABL and LSM

- Favoring the energy conservation: Between two calls of the radiation code, the change of  $T_s$  is supposed not impact  $LW_{net}$   $LW_{net}$  is computed by the radiation code in LMDZ, the radiation in enerbil and in LMDZ is consistent and the radiation is computed in LMDZ.
- $LW_{up}$  is updated as  $T_s$  is updated. The radiation used by enerbil and by LMDZ are not consistent
- $LW_{up}$  can be averaged over the time-steps between two calls of the radiation (account for the variation of  $T_s$  induced by varying surface conditions) (not yet implemented).

# Atmosphere/surface coupling in LMDZOR

LMDZ (phylmd)



# More information on

<http://forge.ipsl.jussieu.fr/orchidee/wiki/Documentation/CouplingLMDZ>