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# The soil carbon in ORCHIDEE

GUENET Bertrand



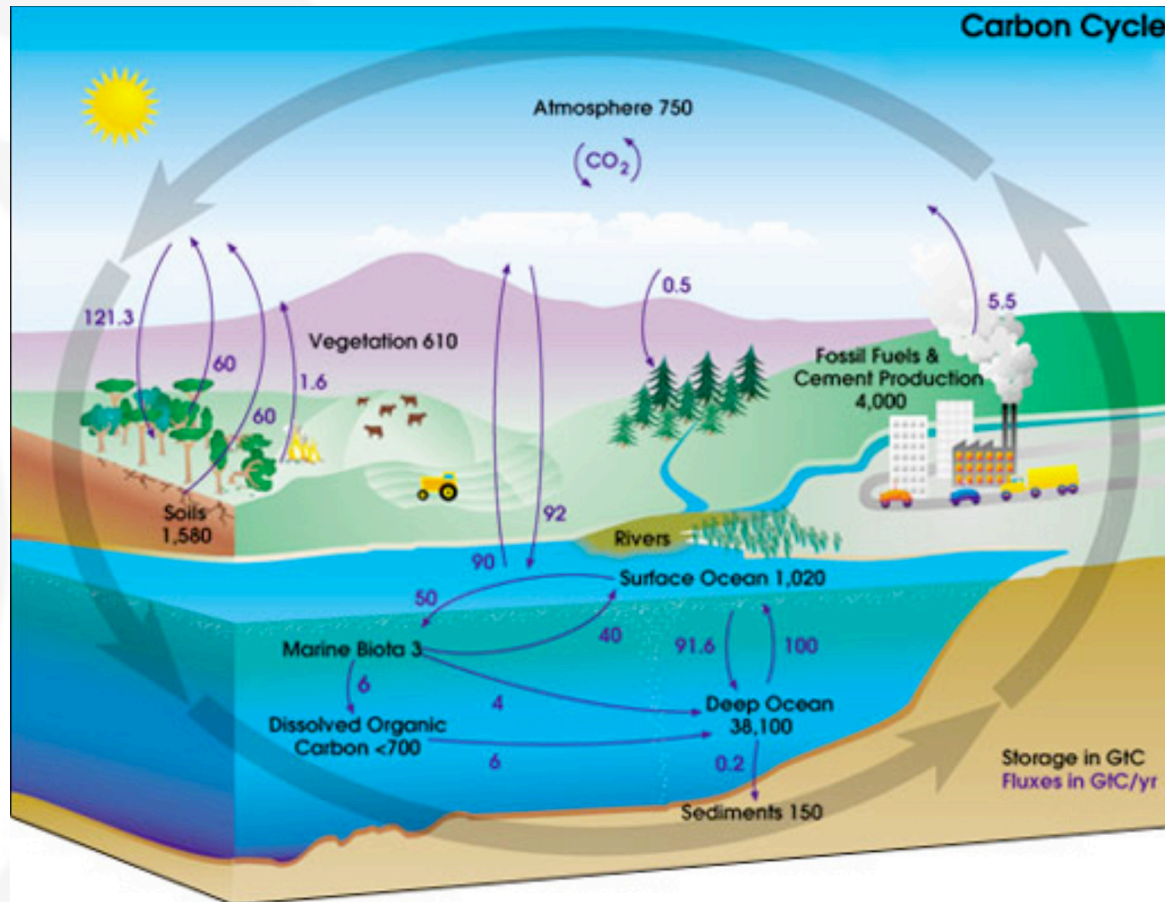
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ORCHIDEE—training Jan. 2020



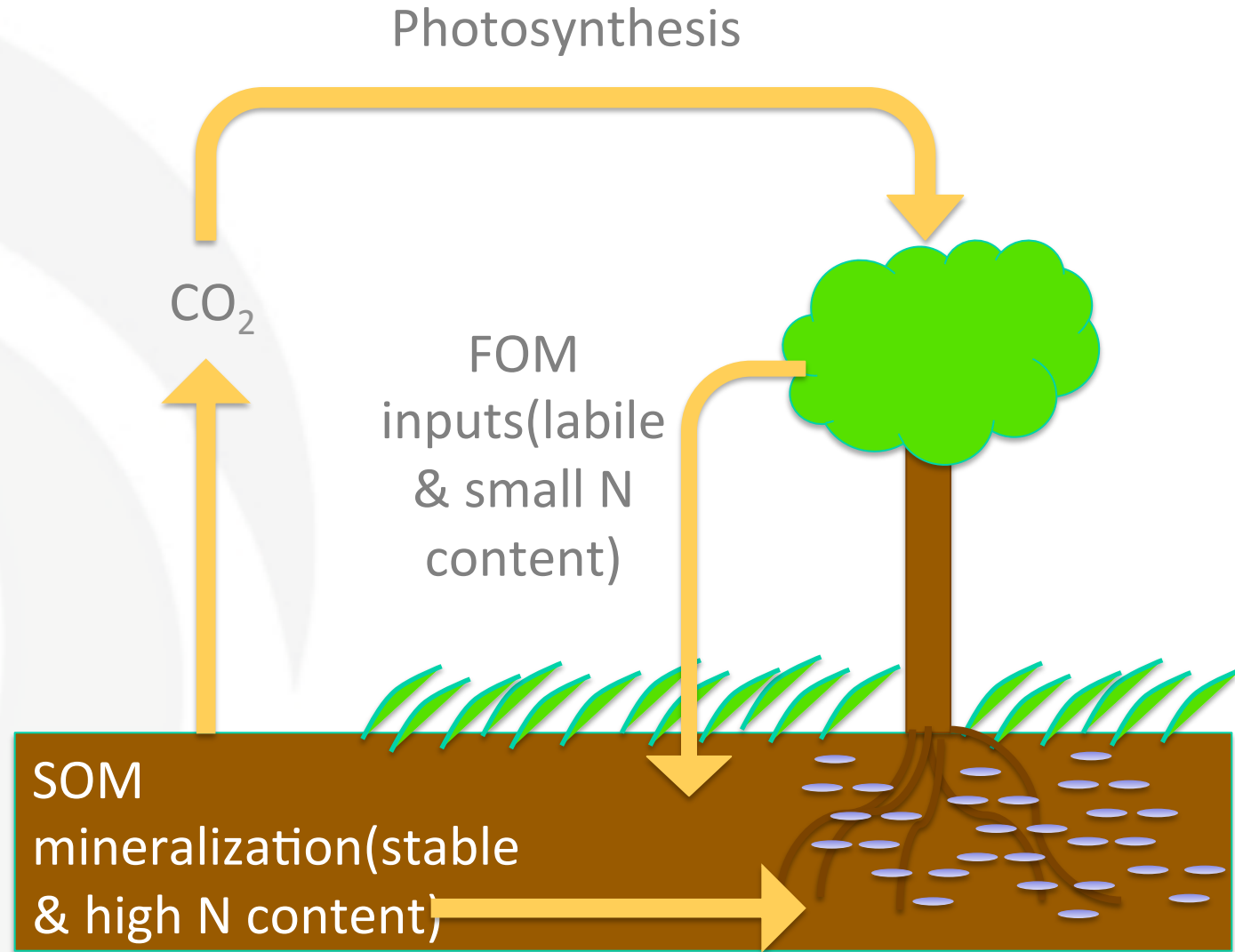
# THE GLOBAL C CYCLE AND ITS SOIL COMPONENT

- The C cycle: a complex cycle composed of different pools.
- These pools interact *via* different fluxes.

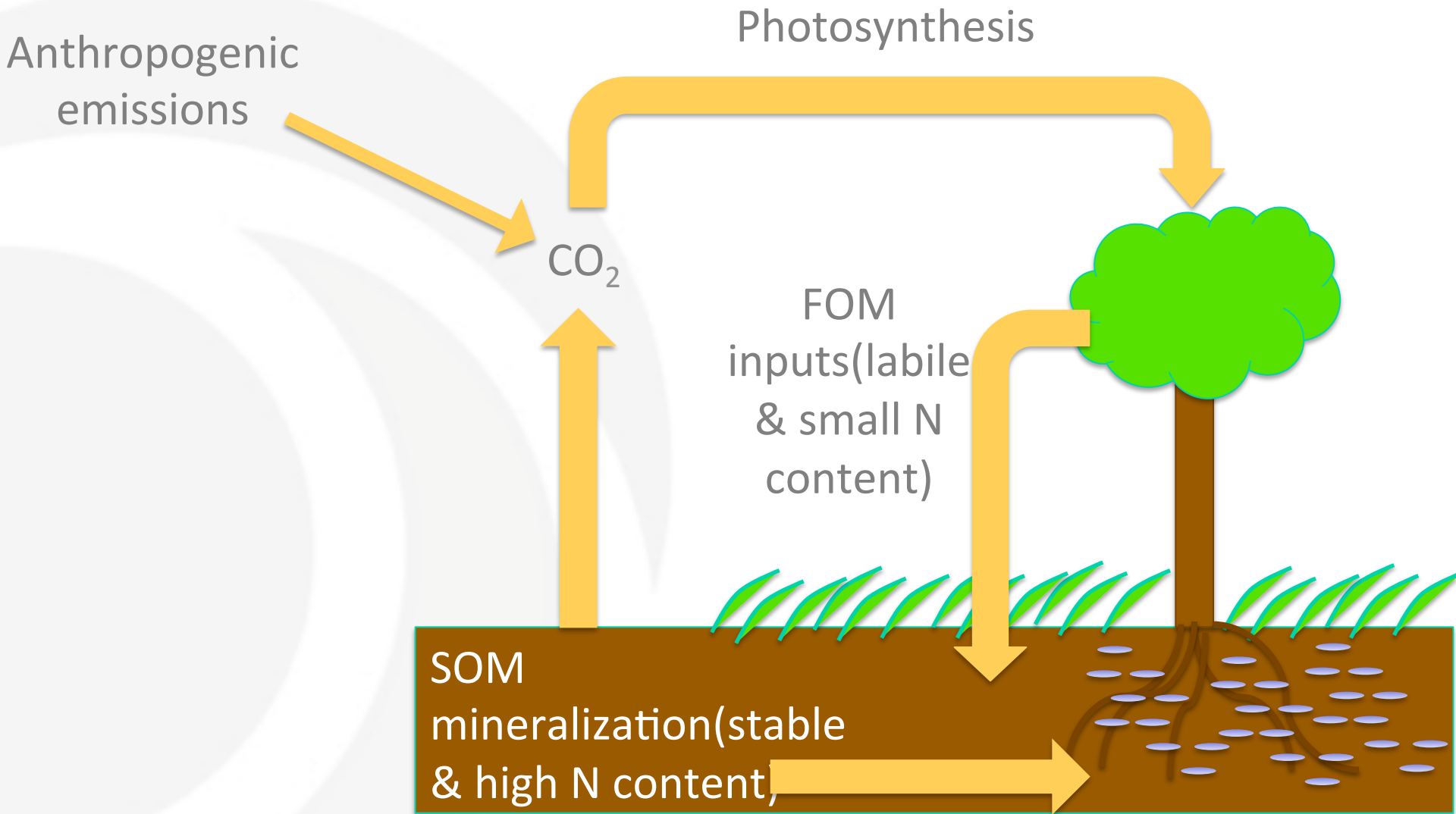


# SOILS: MAJOR ACTORS OF THE C CYCLE

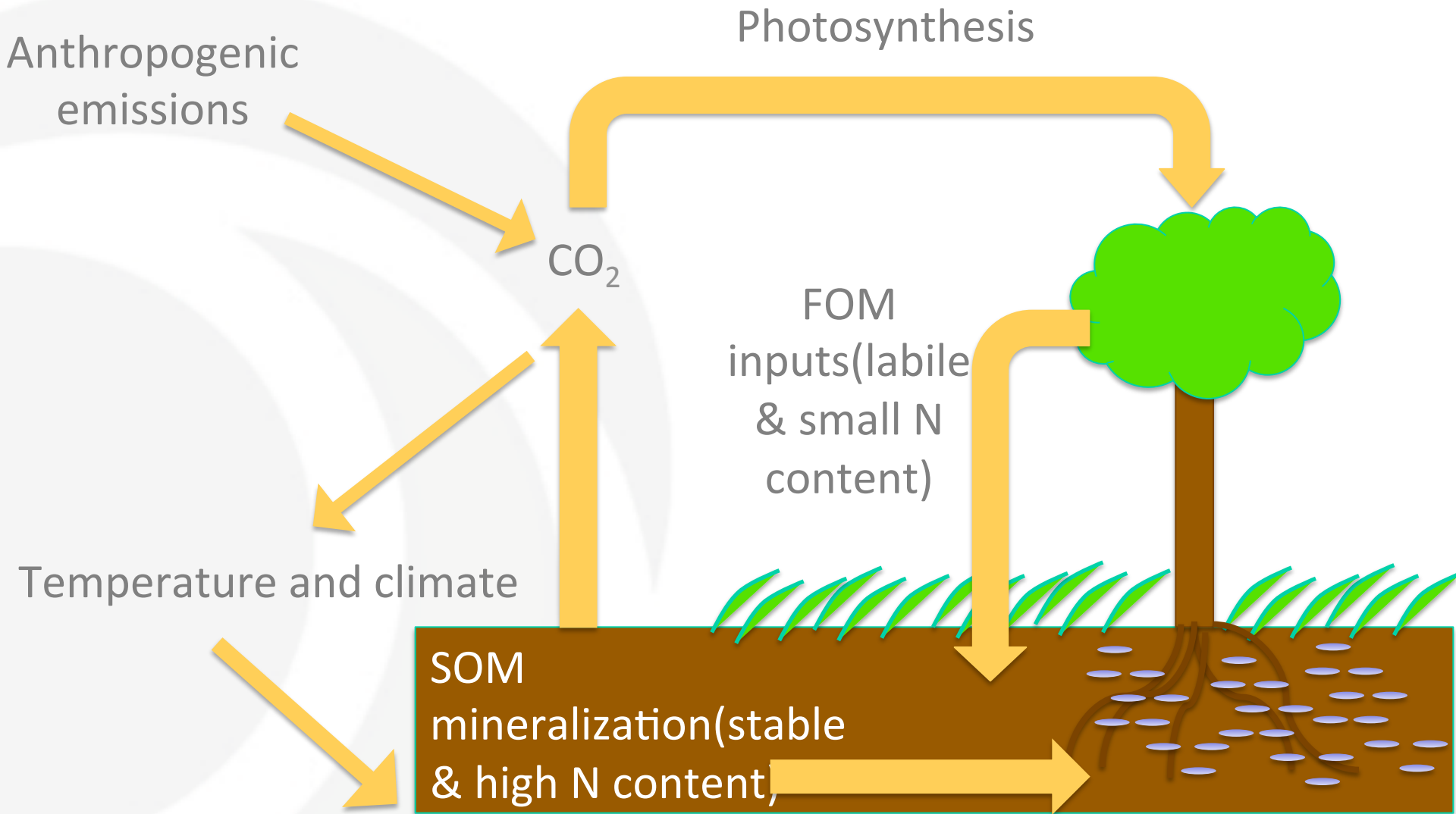
- Temperature and humidity (Reichstein et al., 2003)
- Amount and chemical composition of the Fresh Organic Matter (FOM) (Guenet et al., 2010)



# SOILS AND GLOBAL CHANGES



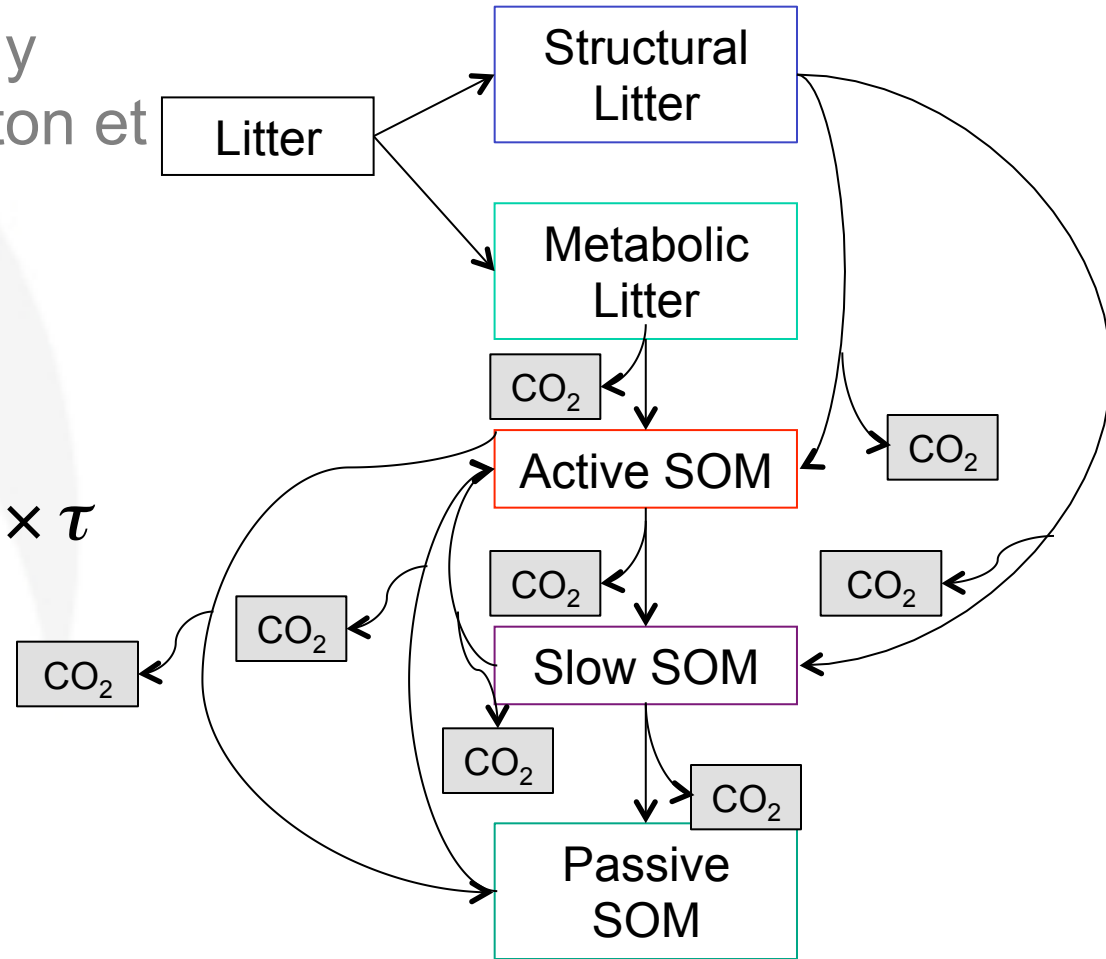
# SOILS AND GLOBAL CHANGES



# THE SOIL C IN ORCHIDEE

- Soil representation mainly based on CENTURY (Parton et al., 1987).

$$\frac{\partial SOC}{\partial t} = I - k \times SOC \times \theta \times \tau$$



# THE SOIL C IN ORCHIDEE

- Split between stomate\_litter.f90 and stomate\_soilcarbon.f90
- Run at ½ hourly time-step whereas stomate runs at daily time-step.
- Moisture and temperature function calculated in stomate\_litter.f90

$$\tau = Q_{10}^{(T - T_{opt})/10}$$

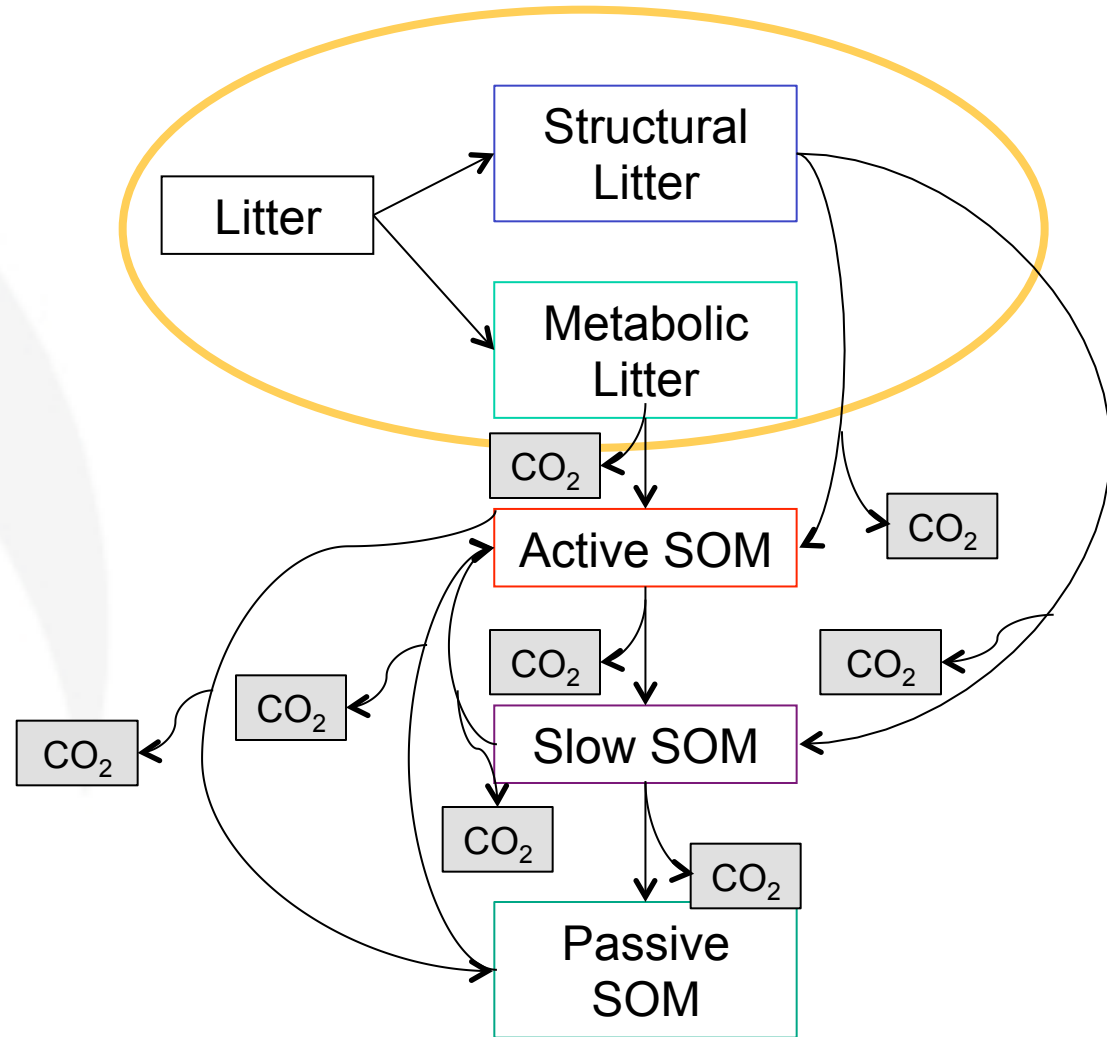
$$\theta = \text{Max}(0.25, \text{Min}(1, M))$$

$$M = -1.1 * SM^2 + 2.4 * SM - 0.29$$



# THE SOIL C IN ORCHIDEE

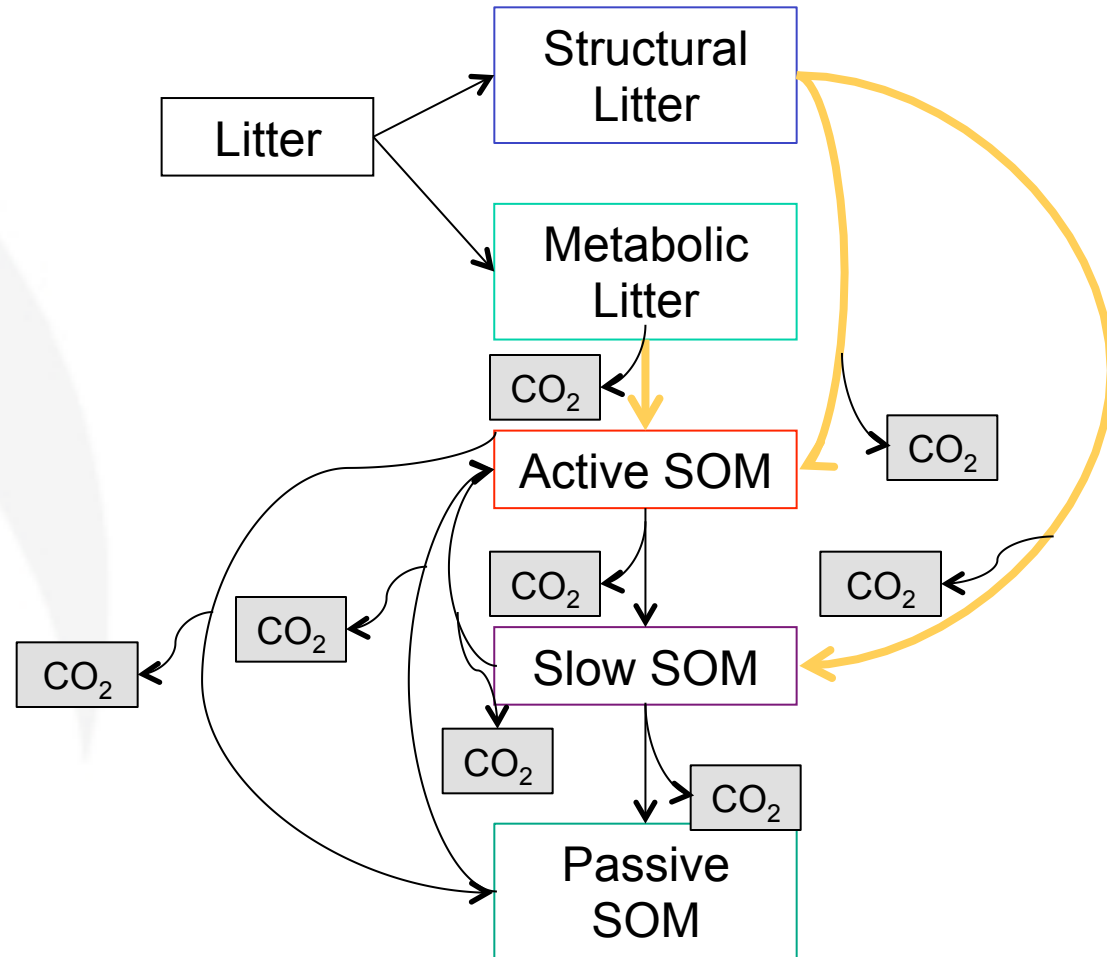
- Input from plants through *bm\_to\_litter* and *turnover*
- Split between above and below ground
- Split into two pools: metabolic/structural depending on lignin and N content of the litter.





# THE SOIL C IN ORCHIDEE

- Inputs from litter decomposition in *soilcarbon\_input*
- Distributed into the active and slow pools control by the lignin content.

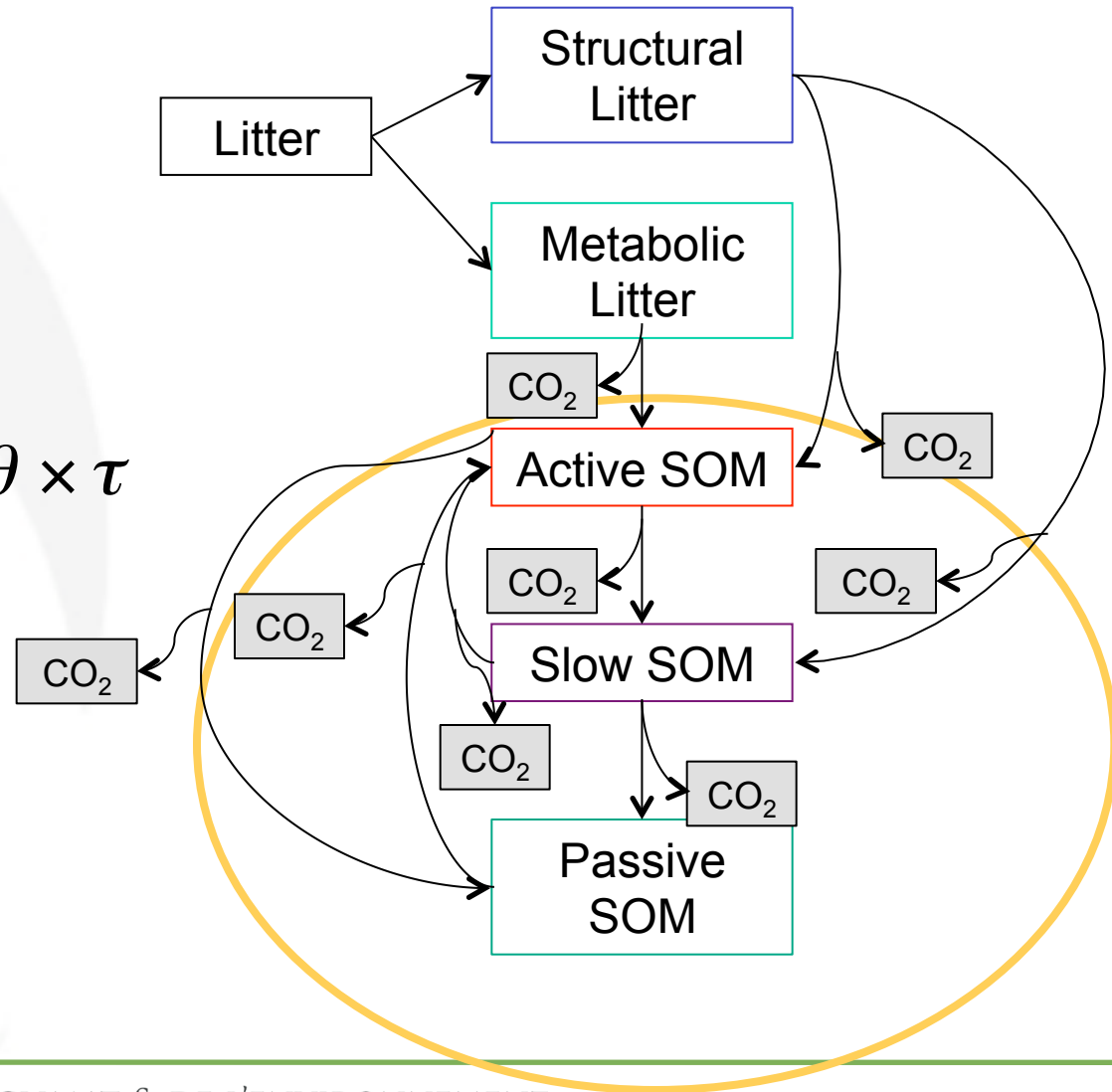


# THE SOIL C IN ORCHIDEE

- Decomposition following 1<sup>st</sup> order kinetics.

$$\frac{\partial SOC}{\partial t} = I - k \times SOC \times \theta \times \tau$$

- A fraction of C decomposed is respired the 1-resp is distributed in the other pools.

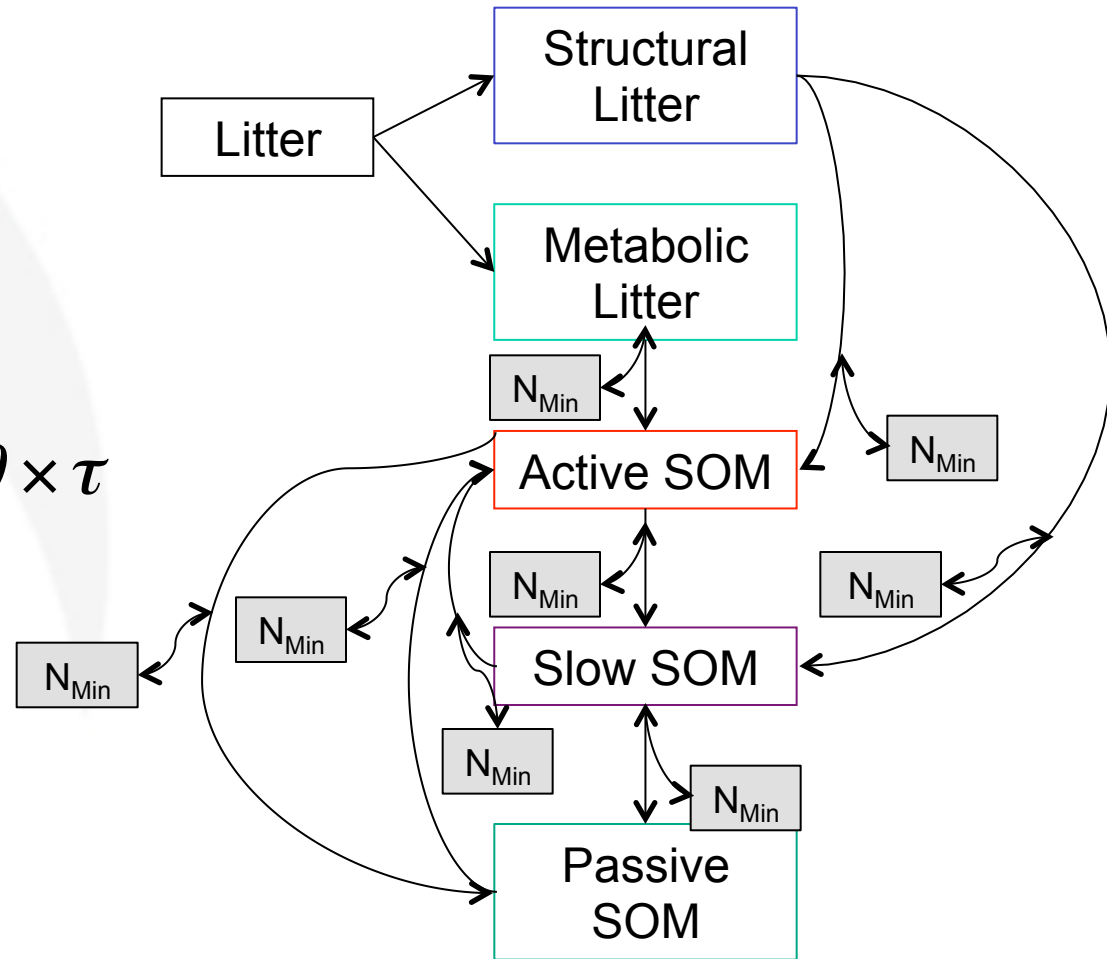


# THE SOIL N IN ORCHIDEE

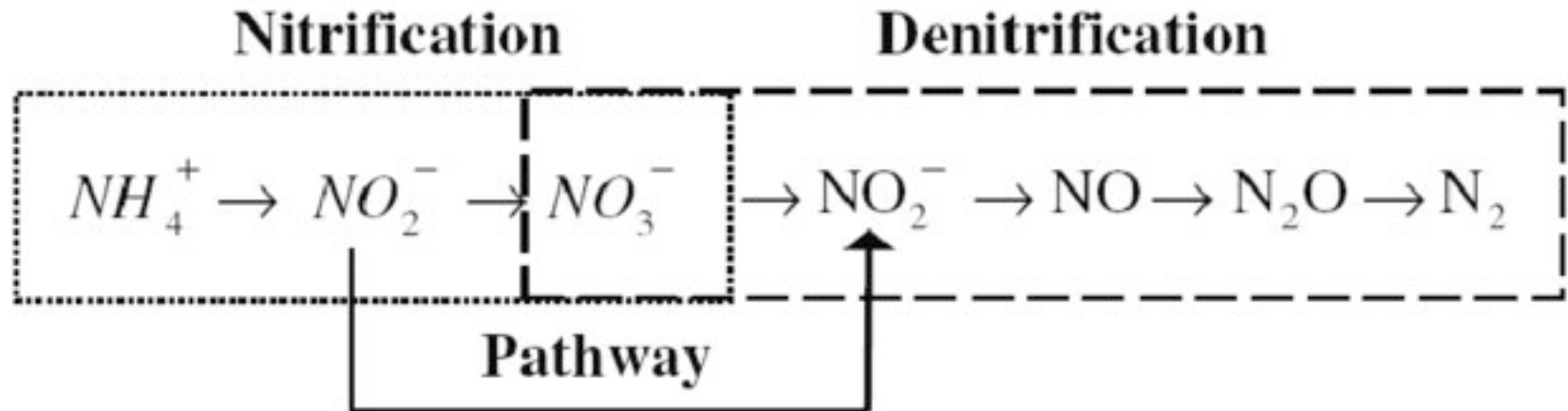
- Decomposition following 1<sup>st</sup> order kinetics.

$$\frac{\partial SON}{\partial t} = I - k \times SON \times \theta \times \tau$$

- The flux of N from one pool to another must satisfy the CN<sub>target</sub> of the receiving pools



# HOW THE NITRIFICATION/DENITRIFICATION PROCESSES ARE REPRESENTED



Peng and Zhu (2006)



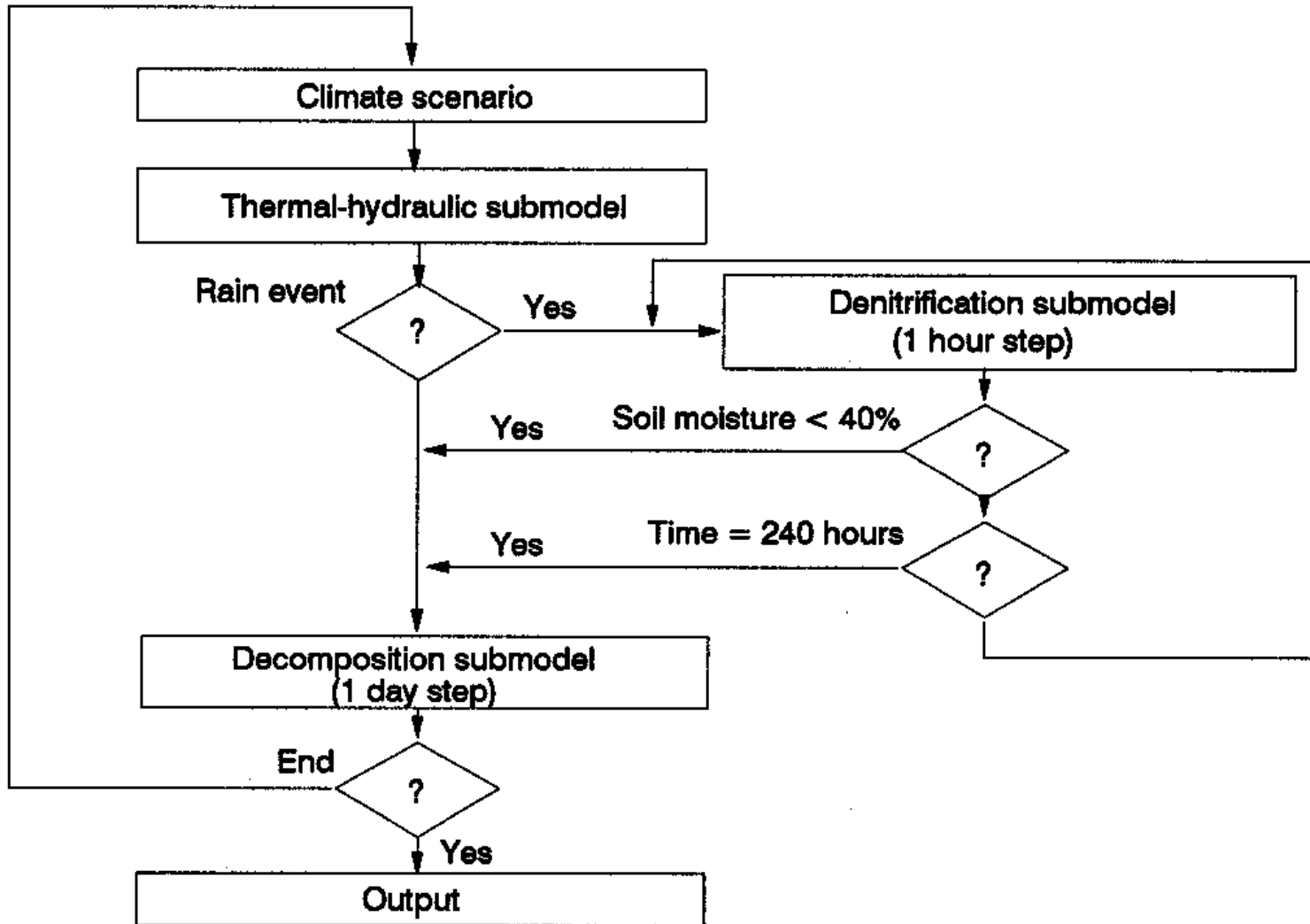
# HOW THE NITRIFICATION/DENITRIFICATION PROCESSES ARE REPRESENTED

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- Key point -> N outputs fluxes & GHG production
- DNDC is an old model based on Li et al. 1992.
- Design to represent denitrification and decomposition.
- In ORCHIDEE, only the N-related aspects are used but in a simplified way.

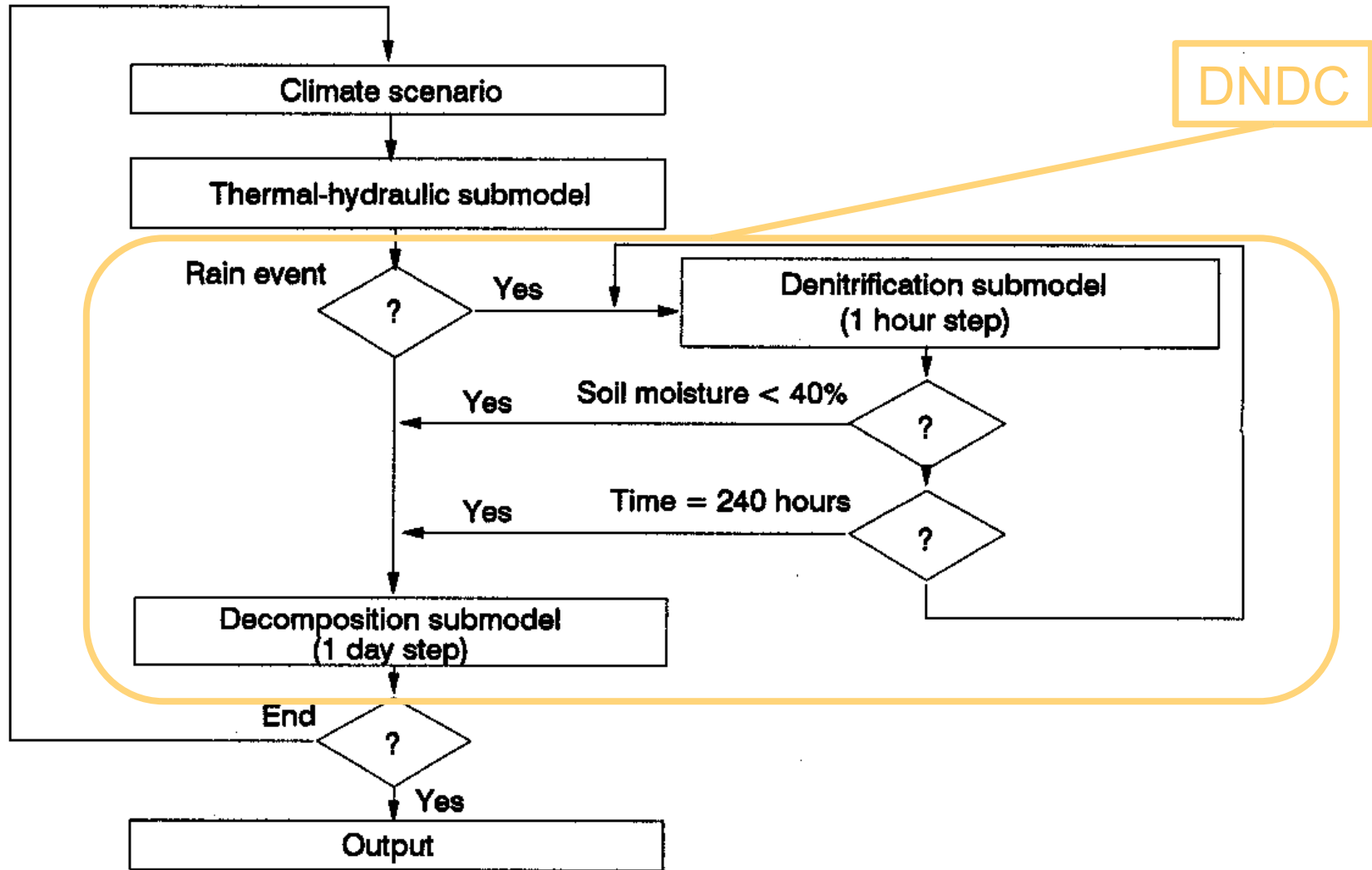


# ORIGINAL DNDC



Li et al., (1992)

# ORIGINAL DNDC

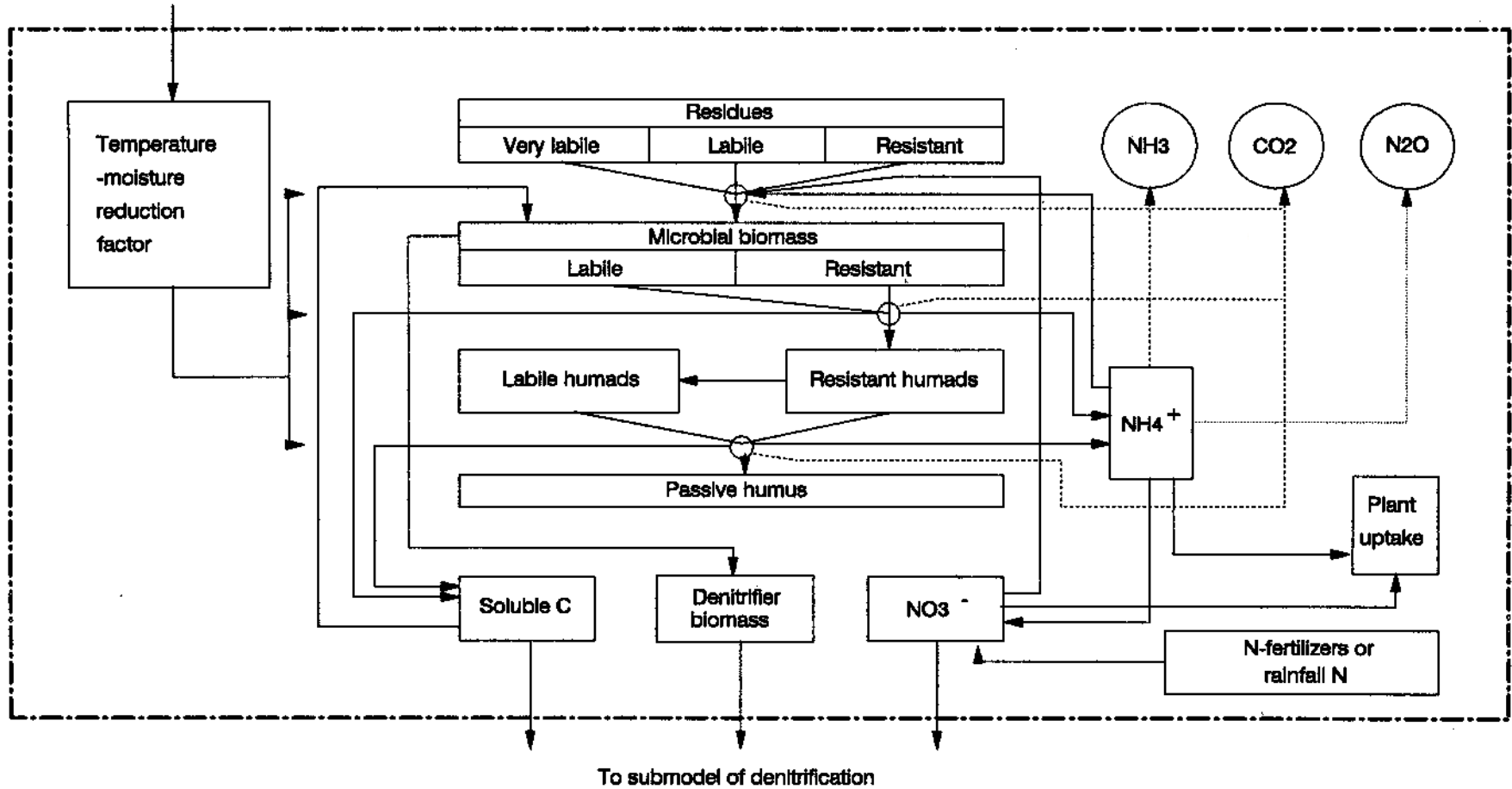


Li et al., (1992)



# ORIGINAL DNDC

Soil temperature & moisture  
from submodel of thermal-  
hydraulic flow



Li et al., (1992)

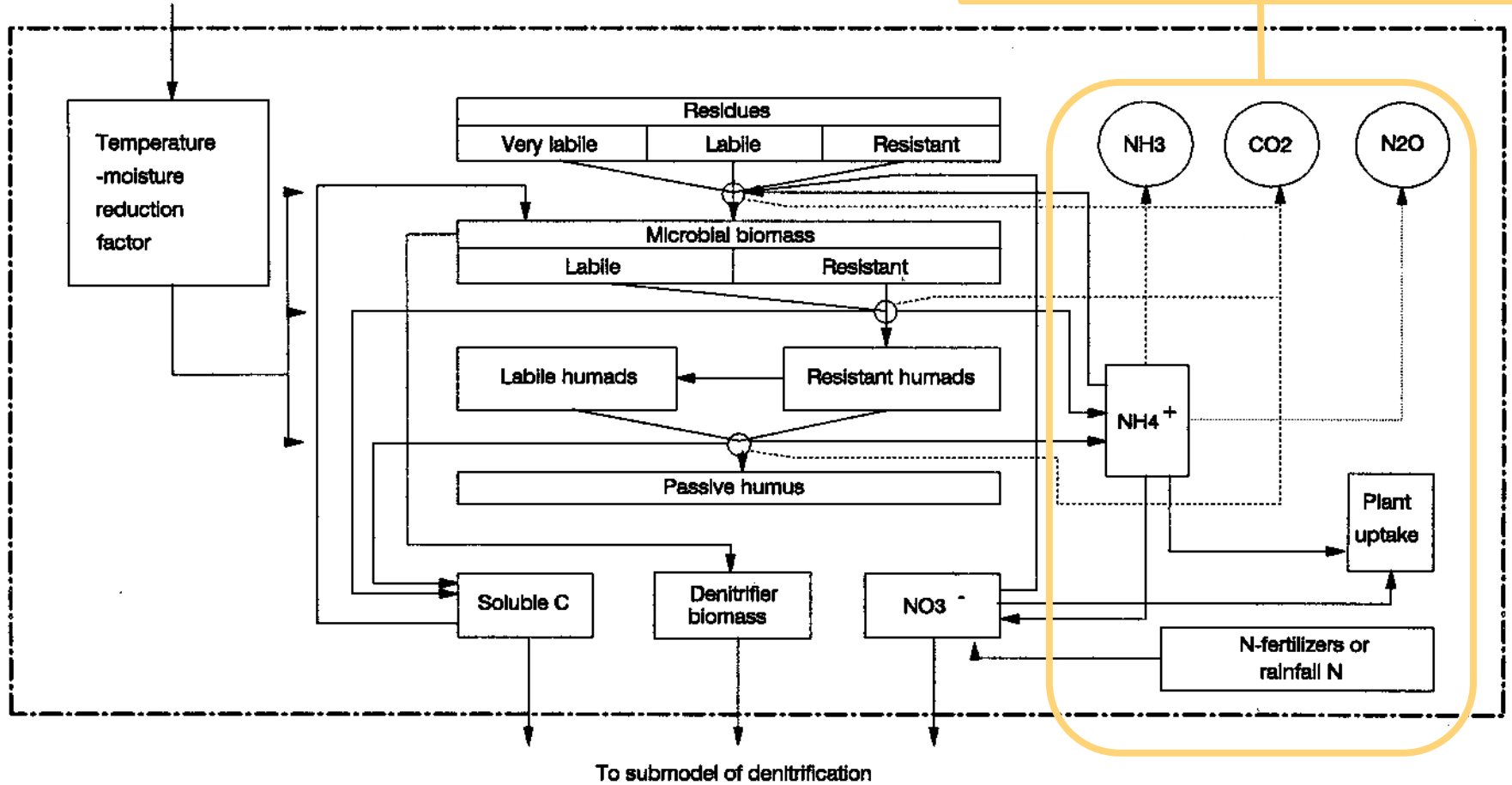




# ORIGINAL DNDC

Soil temperature & moisture  
from submodel of thermal-  
hydraulic flow

USED IN ORCHIDEE

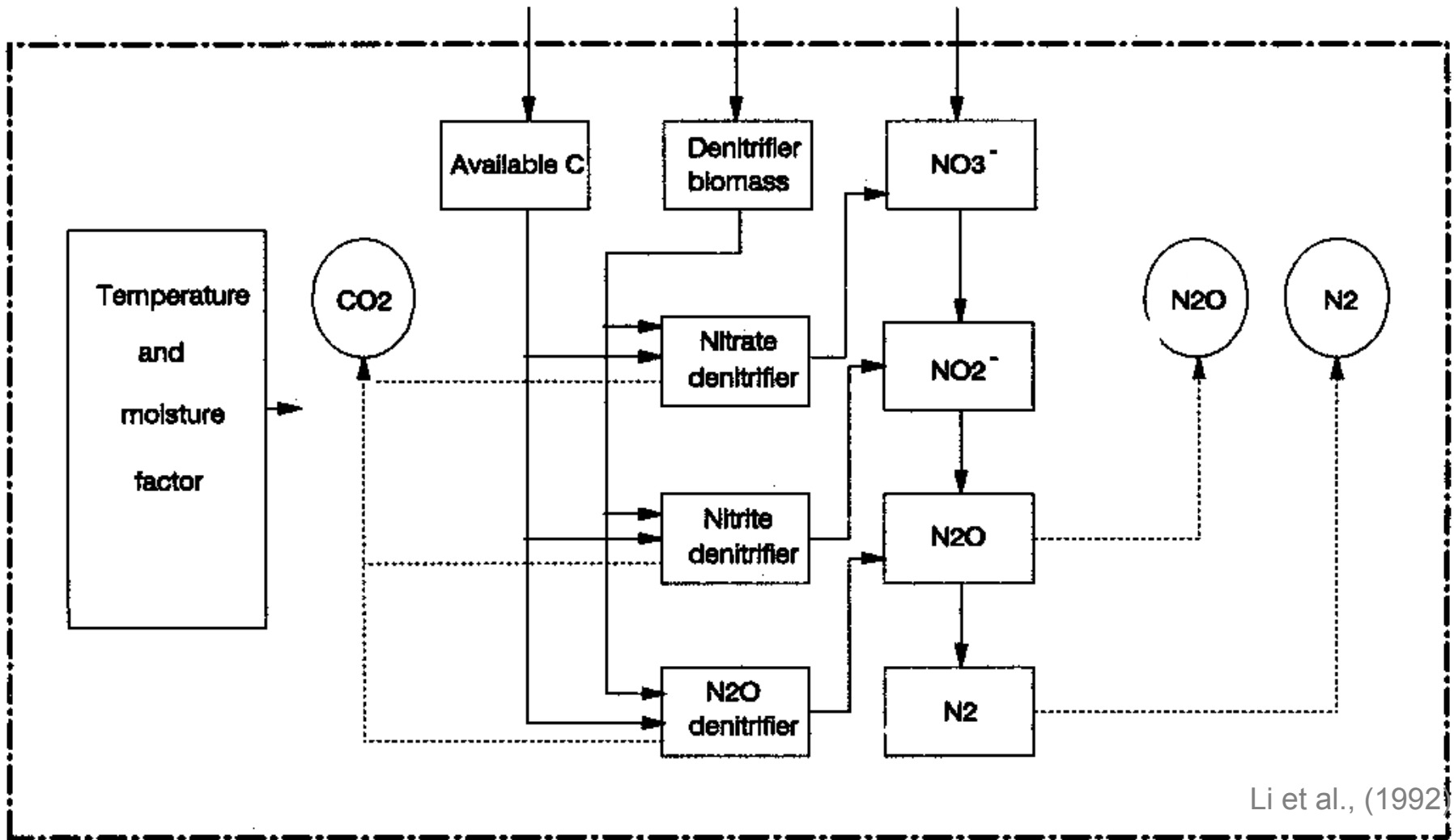


Li et al., (1992)



# ORIGINAL DNDC

From submodel of decomposition

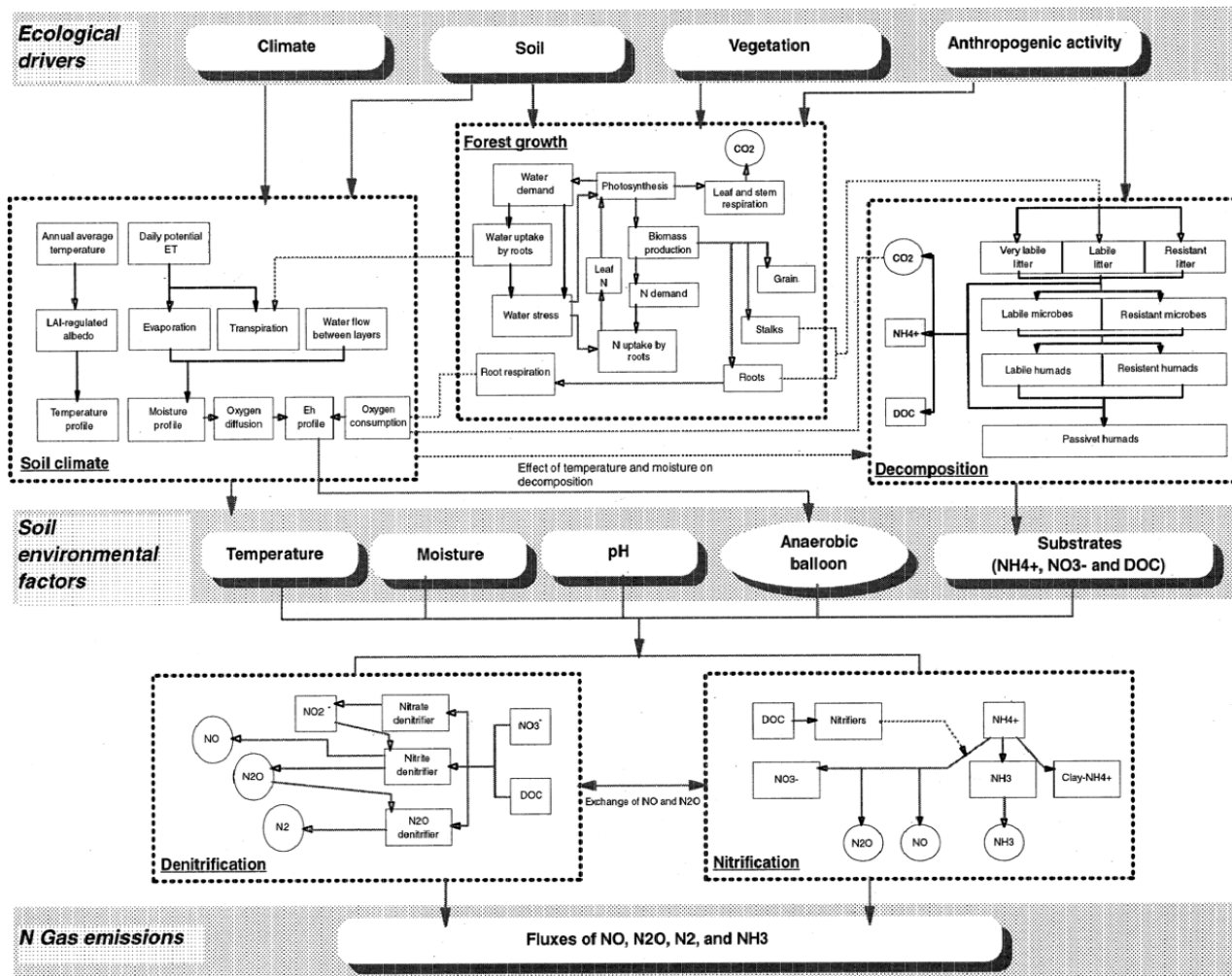


Li et al., (1992)



# IMPROVED DNDC

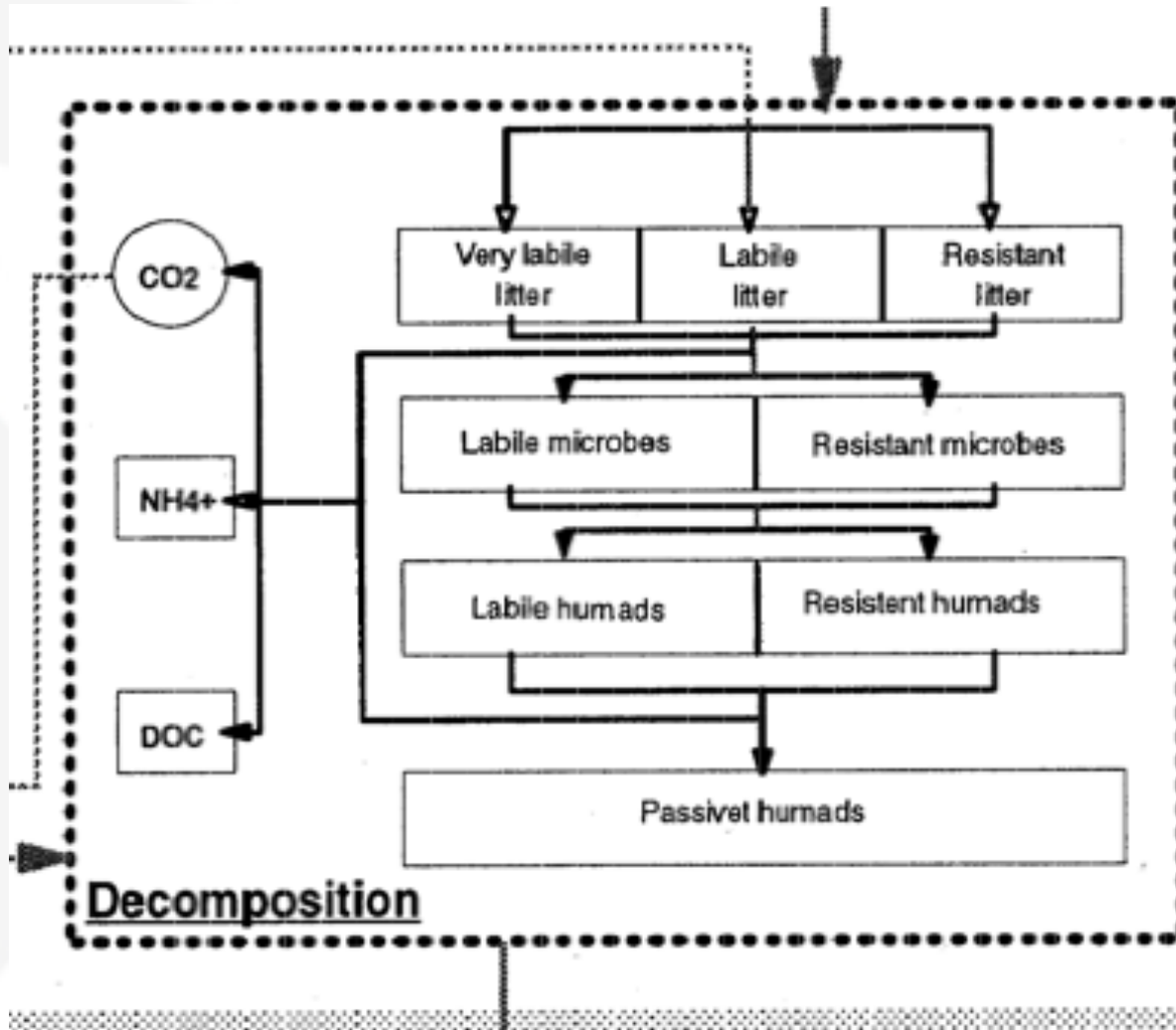
## The PnET-N-DNDC Model



Li et al., (2000)



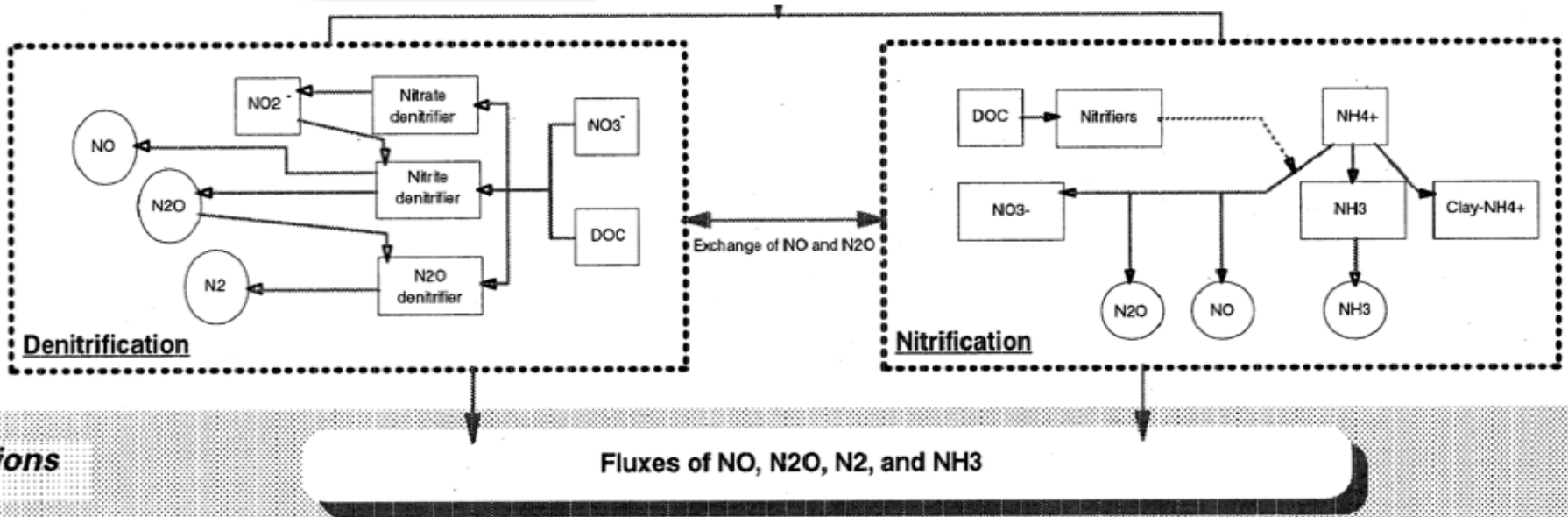
# IMPROVED DNDC



Li et al., (2000)



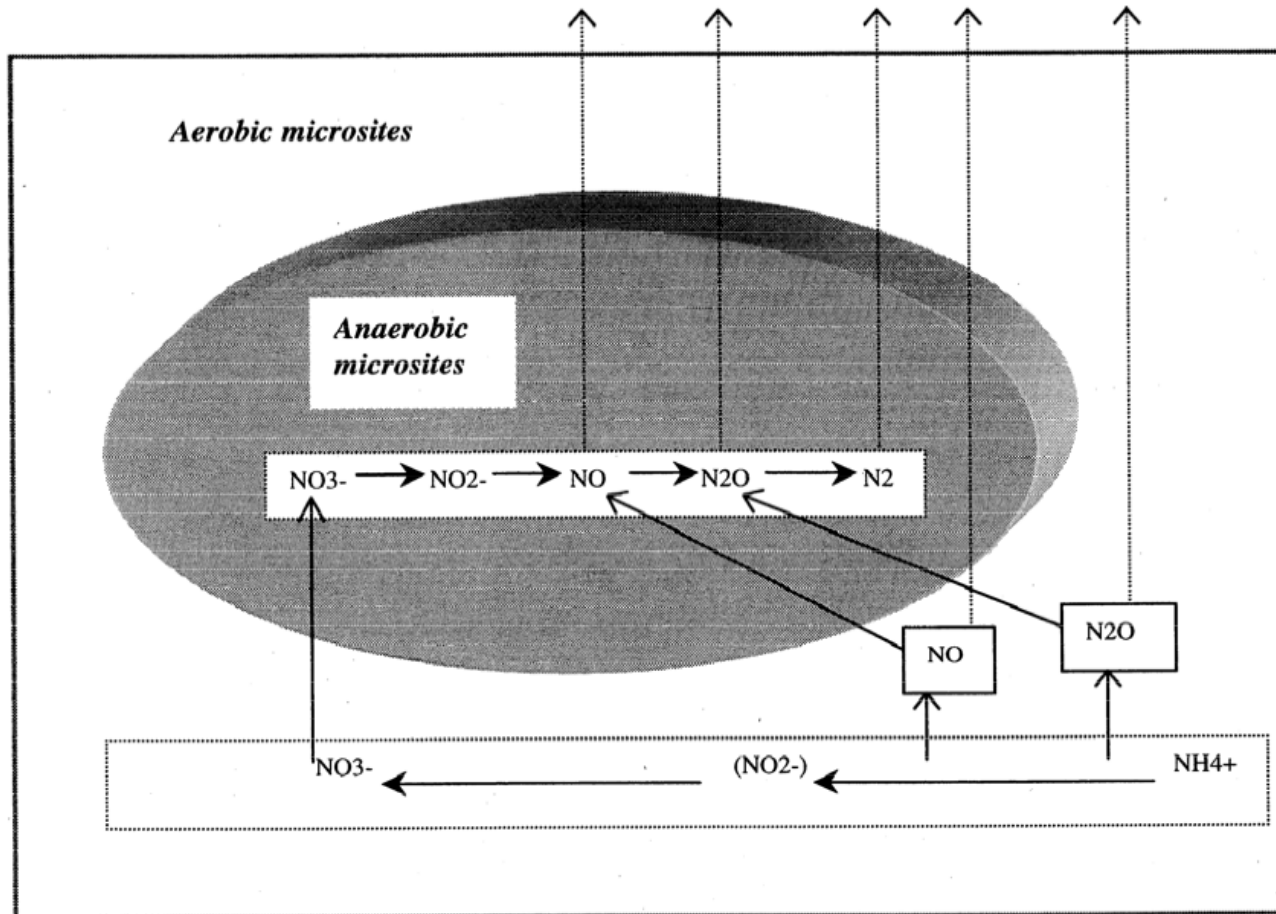
# IMPROVED DNDC



Li et al., (2000)

# THE ANAEROBIC BALLOON CONCEPT

An "Anaerobic Balloon" in Soil Matrix



Li et al., (2000)



# THE ANAEROBIC BALLOON CONCEPT

**Table 2.** Functions and Parameters for O<sub>2</sub> Diffusion and Volumetric Fraction of Anaerobic Microsites (ANVF)

Equation No.	Function	Equation
1	oxygen diffusion coefficient in soil	$D_{s[L]} = D_{air} \text{afps}_{[L]}^{3.33} / \text{afps}_{\text{max}[L]}^{2.0};$
2	oxygen diffusion rate affected by frost	$D_{s[L]} = D_{s[L]} F_{\text{frost}}; 0 < D_{s[L]} < 1$ if $T > 0 \text{ } ^\circ\text{C}, F_{\text{frost}} = 1.2;$ if $T \leq 0 \text{ } ^\circ\text{C}, F_{\text{frost}} = 0.8;$ $D_{s[L]} = D_{s[L]} F_{\text{frost}}; 0 < D_{s[L]} < 1$ if $T > 0 \text{ } ^\circ\text{C}, F_{\text{frost}} = 1.2;$ if $T \leq 0 \text{ } ^\circ\text{C}, F_{\text{frost}} = 0.8;$
3	oxygen partial pressure	$d(pO_{2[L]})/dt = (d(D_{s[L]}) d(pO_{2[L]})/dz)/dz - R)/\text{afps};$
4	Volumetric fraction of anaerobic microsites	$\text{anvf}_{[L]} = a (1 - (b pO_{2[L]} / pO_{2\text{air}}));$

a, b, constant coefficients; afps, air-filled porosity; afps<sub>max</sub>, porosity; anvf, volumetric fraction of anaerobic microsites; D<sub>air</sub>, oxygen diffusion rate in the air, 0.07236 m<sup>2</sup>/h [Beisecker, 1994]; D<sub>s</sub>, oxygen diffusion coefficient in soil; F<sub>frost</sub>, frost factor; L, layer number; pO<sub>2</sub>, oxygen partial pressure; R, oxygen consumption rate (kg C ha<sup>-1</sup>h<sup>-1</sup>); t, time (h); z, soil depth (m).

Li et al., (2000)



# THE EQUATIONS

**Table 3.** Functions and Parameters for Nitrification

Equation No.	Function	Equation
1	relative growth rate of nitrifiers	$\mu_g = \mu_{MAX} ((DOC) / (1 + [DOC]) + F_m / (1 + F_m));$
2	relative death rate of nitrifiers	$\mu_d = a_{MAX} B_n / (5 + [DOC]) / (1 + F_m);$
3	net increase in nitrifiers biomass	$\mu_b = (\mu_g - \mu_d) B_n F_t F_m;$
4	nitrification rate	$R_n = R_{max} [NH4] B_n pH;$
5	temperature factor	$F_t = ((60-T) / 25.78)^{3.503} e^{(3.503 (T-34.22) / 25.78)} ;$
6	moisture factor	if wfps > 0.05 $F_m = 1.01 - 0.21 \text{ wfps};$ if wfps <= 0.05 $F_m = 0;$
7	NO production from nitrification	$NO = .0025 R_n F_t;$
8	N <sub>2</sub> O production from nitrification	$N_2O = 0.0006 R_n F_t \text{ wfps};$

$a_{MAX}$ , maximum death rate for nitrifiers (1.44 1/d [from *Blagodatsky and Richter, 1998*]);  $B_n$ , biomass of nitrifiers (kg C/ha); [DOC], concentration of dissolved organic C (kg C/ha);  $F_m$ , moisture factor;  $F_t$ , temperature factor; [NH4], concentration of ammonium (kg N/ha); NO, NO production from nitrification; N<sub>2</sub>O, N<sub>2</sub>O production from nitrification [*Ingwersen et al., 1999*]; pH, soil pH;  $R_n$ , nitrification rate;  $R_{max}$ , maximum nitrification rate (1/h); T, soil temperature (°C); wfps, water-filled porosity;  $\mu_{MAX}$ , maximum growth rate for nitrifiers (4.87 1/d [from *Blagodatsky and Richter, 1998*]);  $\mu_b$ , net increase in nitrifiers biomass;  $\mu_d$ , relative death rate of nitrifiers;  $\mu_g$ , relative growth rate of nitrifiers.

Li et al., (2000)





# THE EQUATIONS

**Table 4.** Functions and Parameters for Denitrification

Equation No.	Function	Equation
1	relative growth rate of No <sub>x</sub> denitrifiers	$\mu_{NO_x} = \mu_{NO_x(max)} [DOC]/(K_c + [DOC]) [No_x]/(K_n + [NO_x]);$
2	relative growth rate of total denitrifiers	$\mu_g = F_t (\mu_{NO_3} F_{PH1} + \mu_{NO_2} F_{PH2} + \mu_{NO} F_{PH2} + \mu_{N_2O} F_{PH3});$ $F_t = 2^{((T-22.5)/10)};$ $F_{PH1} = 1 - 1 / (1 + e^{[pH-4.25/0.5]});$ $F_{PH2} = 1 - 1 / (1 + e^{[pH-5.25/1.0]});$ $F_{PH3} = 1 - 1 / (1 + e^{[pH-6.25/1.5]});$
3	denitrifier growth rate, death rate, and consumption rate of soluble carbon	$R_g = \mu_g B_d;$ $R_d = M_c Y_c B_d;$ $R_C = (\mu_g / Y_c + M_c) B_d;$
4	consumption rates of N oxides	$R_{NO_x} = (\mu_{NO_x} / Y_{NO_x} + M_{NO_x} [No_x]/[N]) B_d;$
5	nitrogen assimilation rate	$q_N = R_g / CN;$
6	gas diffusion factor	$v = D_{max} afps (1 - anv) F_{clay} 2^{T/20};$ $F_{clay} = 0.13 - 0.079 \text{ clay};$

afps, air-filled porosity; anv, volumetric fraction of anaerobic microsites;  $B_d$ , denitrifier biomass (kg C/m<sup>3</sup>); clay, clay fraction in the soil; CN, C/N ratio in denitrifiers (3.45 [Van Verseveld and Stouthamer 1978]);  $D_c$ , consumption rate of soluble carbon by denitrifiers (kg C m<sup>-3</sup>h<sup>-1</sup>);  $D_{max}$ , maximum diffusion rate in air (m<sup>2</sup>/h);  $D_{NO_x}$ , consumption rate of N oxides by denitrifiers (kg C m<sup>-3</sup>h<sup>-1</sup>); [DOC], soluble C concentration (kg C/m<sup>3</sup>);  $F_{clay}$ , clay factor;  $F_t$ , temperature factor;  $F_{PH1}$ , pH factors for NO<sub>3</sub><sup>-</sup> denitrifiers;  $F_{PH2}$ , pH factors for NO<sub>2</sub><sup>-</sup> and NO denitrifiers;  $F_{PH3}$ , pH factors for N<sub>2</sub>O denitrifiers;  $K_c$ , half-saturation value of soluble carbon (0.017 kg C/m<sup>3</sup> [Shan and Coulman, 1978]);  $K_n$ , half-saturation value of N oxides (0.083 kg N/m<sup>3</sup> [Shan and Coulman, 1978]);  $M_c$ , maintenance coefficient on carbon (0.0076 kg N kg<sup>-1</sup>h<sup>-1</sup> [Van Verseveld et al., 1977]); [N], concentration of all NO<sub>x</sub> (kg N/m<sup>3</sup>); [No<sub>x</sub>], concentration of for NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup>, NO and N<sub>2</sub>O (kg N/m<sup>3</sup>); pH, soil pH;  $q_n$ , nitrogen assimilation rate (kg N ha<sup>-1</sup>h<sup>-1</sup>); T, soil temperature (°C); v, gas diffusion factor (%);  $Y_c$ , maximum growth rate of denitrifiers on soluble carbon (0.503 kg C/kg C [Van Verseveld et al., 1977]);  $M_{NO_x}$ , maintenance coefficient on N oxides (0.09, 0.035 and 0.079 kg N/kg/ for NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup> (+NO\*) and N<sub>2</sub>O, respectively, based on Van Verseveld et al. [1977]);  $R_d$ , denitrifier death rate;  $R_g$ , denitrifier growth rate;  $Y_{NO_x}$ , maximum growth rate on N oxides (0.401, 0.428 and 0.151 kg C/kg N for NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup> (+NO\*) and N<sub>2</sub>O, respectively, based on Van Verseveld et al. [1977]);  $\mu_g$ , relative growth rate of total denitrifiers (1/h);  $\mu_{NO_3}$ ,  $\mu_{NO_2}$ ,  $\mu_{NO}$ ,  $\mu_{N_2O}$ , relative growth rate of NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup>, NO-, and N<sub>2</sub>O denitrifiers;  $\mu_{NO_x}$ , relative growth rate of NO<sub>x</sub> denitrifiers (1/h);  $\mu_{NO_x(max)}$ , maximum growth rates (0.67 1/h for NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup> denitrifiers, and 0.34 1/h for NO and N<sub>2</sub>O denitrifiers, based on Hartel and Alexander [1987]). The parameters are shared by NO<sub>2</sub><sup>-</sup> and NO due to the lack of data for NO.

Li et al., (2000)



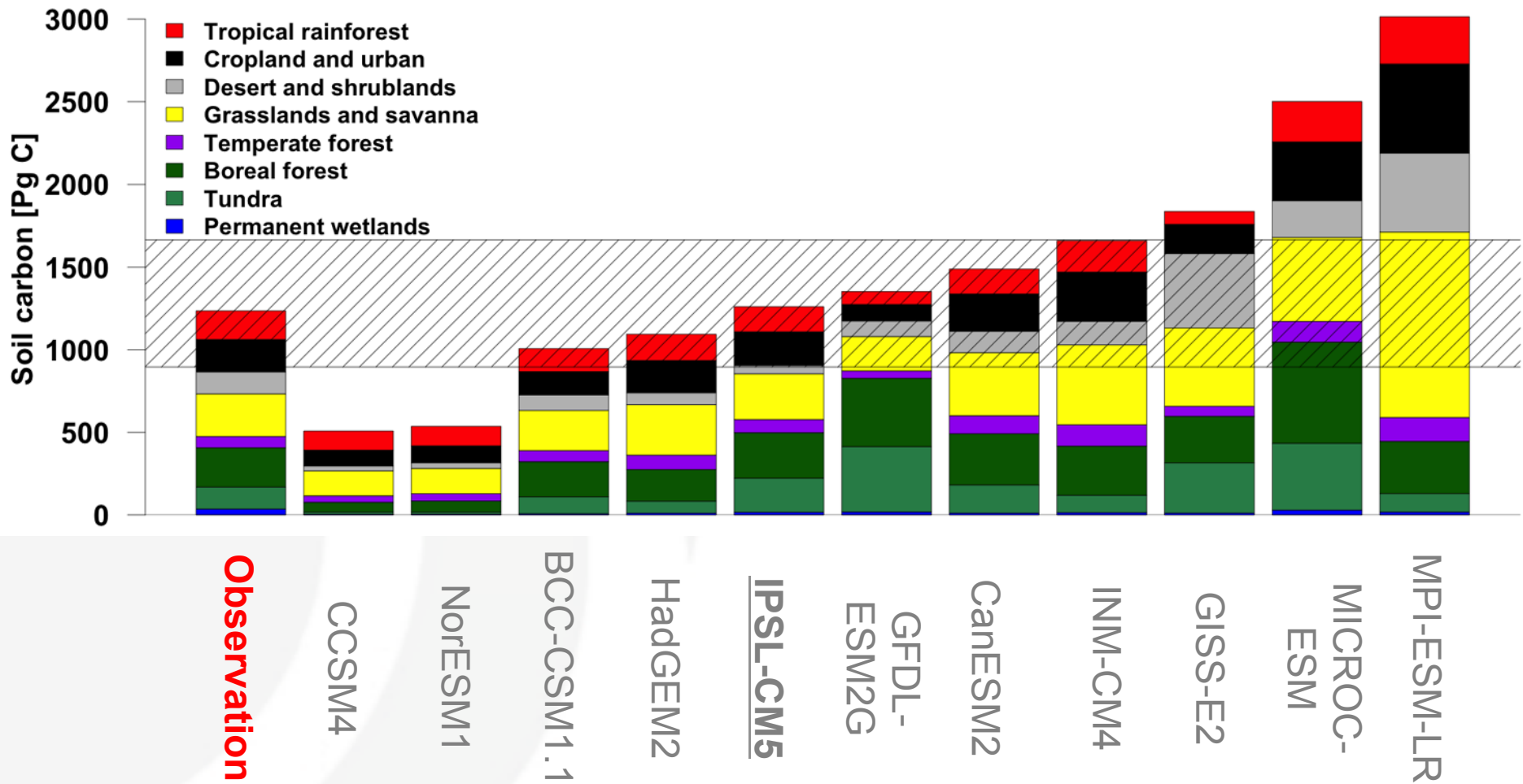
# WHAT IS DONE IN ORCHIDEE

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- Implemented by Zaehle et al., (2010)
- At that time, no DOC and no soil C discretization.
- Active C pools used instead of DOC.
- Gas diffusion is calculated using a fixed soil depth value of 20cm.
- Several parameters tuned.
- With Nicolas, we decided to let the original parameter values.



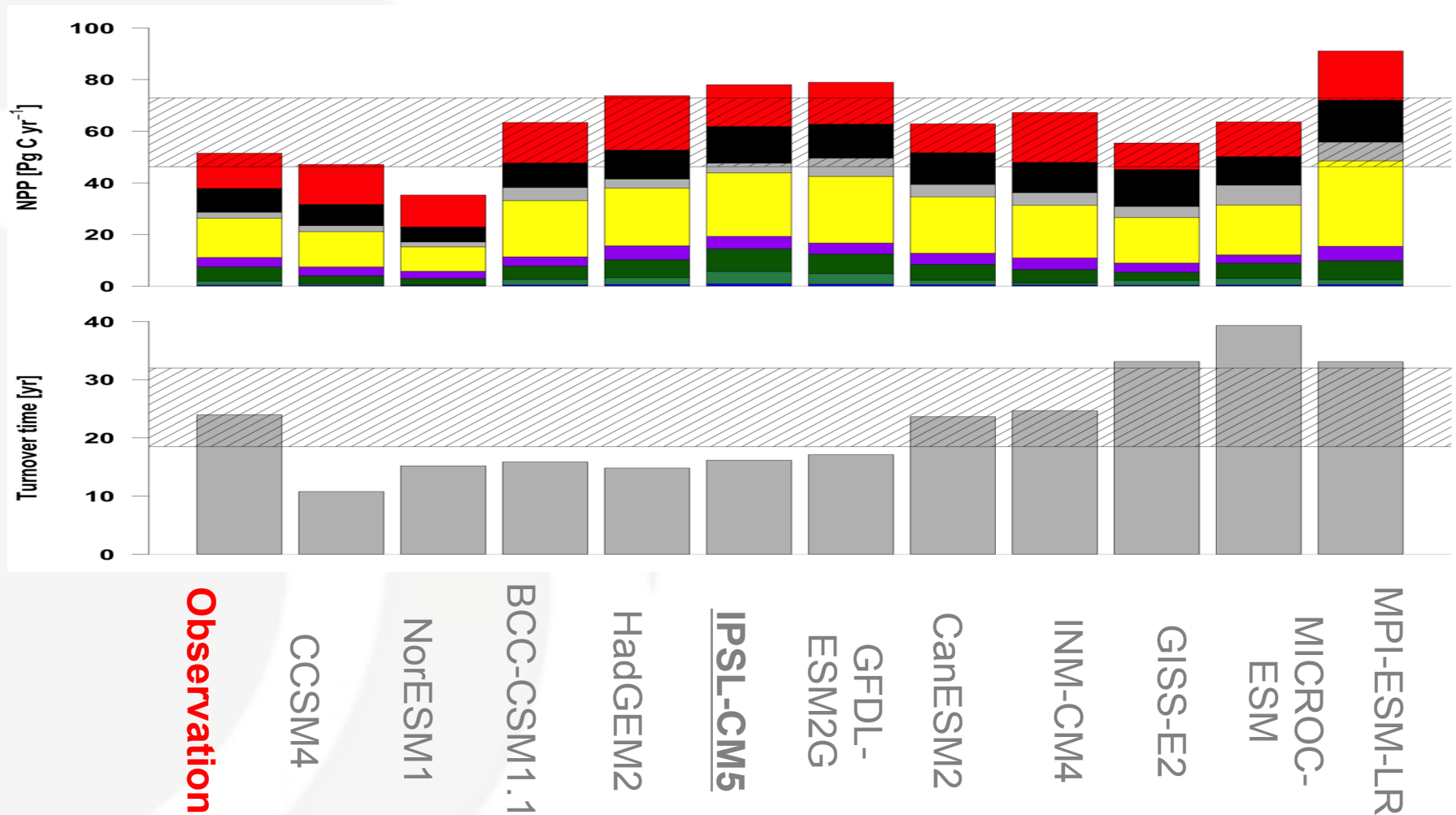
# HOW GOOD ARE EARTH SYSTEM MODELS TO REPRESENT SOIL C STOCK



Todd-Brown et al. (2013)



# HOW GOOD ARE EARTH SYSTEM MODELS TO REPRESENT SOIL C STOCK



Todd-Brown et al. (2013)



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ORCHIDEE—training Jan. 2020



# MEAN RESIDENCE TIME

- Based on  $^{14}\text{C}$  data, ESM MRT are overestimated by ~40%

**Table 1. Global soil carbon stocks and carbon uptake for CMIP5 models that experienced a quadrupling of atmospheric  $\text{CO}_2$  from a preindustrial value of 285 ppm over a period of 140 years.**

ESM	Initial SOC (Pg C)	% change in SOC	% change in SOC after $^{14}\text{C}$ con-straint	$^{14}\text{C}$ -imposed sink reduction (%)	$\tau_{\text{slow}}$ (year)*	$\tau_{\text{passive}}$ (year)	$r_f$	$r_s$	$^{14}\text{C}$ -imposed correction factors <sup>†</sup>			
									$\tau_{\text{slow}}$	$\tau_{\text{passive}}$	$r_f$	$r_s$
CESM1 (BGC)	571	6.3	5.1	19	56 ± 16	1310 ± 241	0.06 ± 0.05	0.33 ± 0.05	–	3.7 ± 1.5	–	0.34 ± 0.75
GFDL-ESM2M	1344	26	3.3	87	231 ± 196	–	0.17 ± 0.07	–	16 ± 18	–	0.06 ± 0.14	–
HadGE M2-ES	1028	63	33	46	208 ± 84	–	0.12 ± 0.07	–	17 ± 12	–	0.07 ± 0.32	–
IPSL-CM5A-LR	1340	27	25	5.9	218 ± 82	1181 ± 347	0.06 ± 0.03	0.29 ± 0.07	–	14 ± 8.3	–	0.07 ± 0.14
MRI-ESM1 <sup>‡</sup>	1403	36	22	40	347 ± 117	1065 ± 257	0.17 ± 0.09	0.10 ± 0.06	–	13 ± 7.2	0.46 ± 0.79	0.34 ± 0.74
Mean <sup>§</sup>	1137 ± 312	32 ± 18	18 ± 12	40 ± 27	212 ± 104	1185 ± 123	0.12 ± 0.06	0.24 ± 0.12	16.5 ± 0.5	10.2 ± 4.6	–	–

\* $\tau_{\text{slow}}$ ,  $\tau_{\text{passive}}$  denote the turnover time, and  $r_f$ ,  $r_s$  denote the transfer coefficient from the fast to the slow pool and from the slow to the passive pool, respectively. Reported values were estimated as an area-weighted mean and standard deviation of all model grid cells. †The mean and standard deviation of the  $^{14}\text{C}$ -imposed correction factors were derived from using the  $^{14}\text{C}$  observations at each site in a single optimization and then averaging these scalar adjustments across the set of 157 optimizations. ‡The  $^{14}\text{C}$ -constrained sink reduction and correction factor for MRI were based on an inverse analysis that changed the pool size of both slow and passive pools. The reported percentage change in SOC and sink reduction were derived from transient simulations starting at steady state with the reduced complexity model. See methods in the supporting materials. §The multimodel mean and standard deviation were estimated using the mean value from each of the five ESMs.

He et al., 2016



# A NEW SOIL SCHEME FROM THE MICT BRANCH

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- For CMIP6, some LSMs will have permafrost C, explicit N cycle and perhaps both
- All the necessary piece of code exist within the « orchidee environment »
- Opportunity to benefit from the huge effort done in the MICT branch by many colleagues (Dan Zhu, Philippe Ciais, Matthieu Guimberteau, Charlie Koven, ...)

**Only an option in the trunk controlled by  
OK\_SOIL\_CARBON\_DISCRETIZATION**



# WHAT DO WE TOOK FROM MICT?

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- Several options are available in MICT but not in the trunk (fire, grassland management, permafrost C)
- Focus on permafrost C
- Not only adding « frozen C »
- Soil C is discretized
- Diffusion is added (including bioturbation and cryoturbation)
- Temperature effect on SOC mineralization
- When frozen, nroot is set to -> impact on water stress and on transpiration.
- Optional
  - Zimov effect
  - Insolation effect (thermal conductivity affected by SOC)



# WHAT IS NEW

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- Soil organic N is also discretized
- Mineral N is not
- No effect on N plant uptake





# ONGOING DEVELOPEMENT

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- Representation of the soil C/N profile
- Lateral outputs of C (DOC, Erosion)
- Representation of Priming effect.
- Carbon isotopes ( $^{14}\text{C}$  and  $^{13}\text{C}$ )
- Peatlands



# Soil Carbon discretization

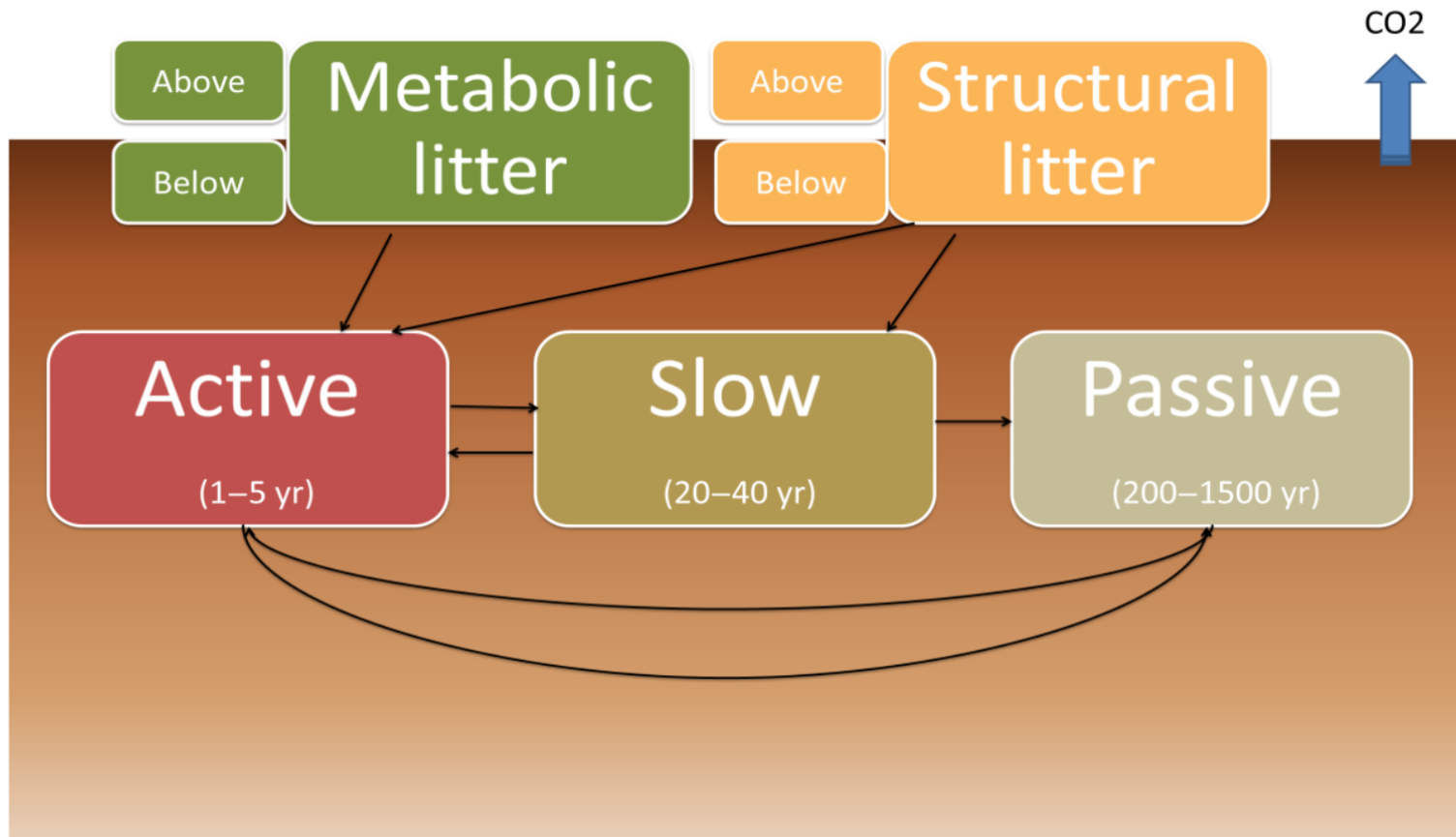
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- Any models used for CMIP5 represent the soil C profiles.
- A substantial part of the soil C stored in deep layers (Jobbagy and Jackson, 2000)
- Deep C dynamic different from surface C (Fontaine et al., 2007)
- In ORCHIDEE any C is lost by drainage or runoff instead of the importance of allochthonous C in the aquatic ecosystems functioning (Cole et al., 2007, Bianchi et al., 2011)



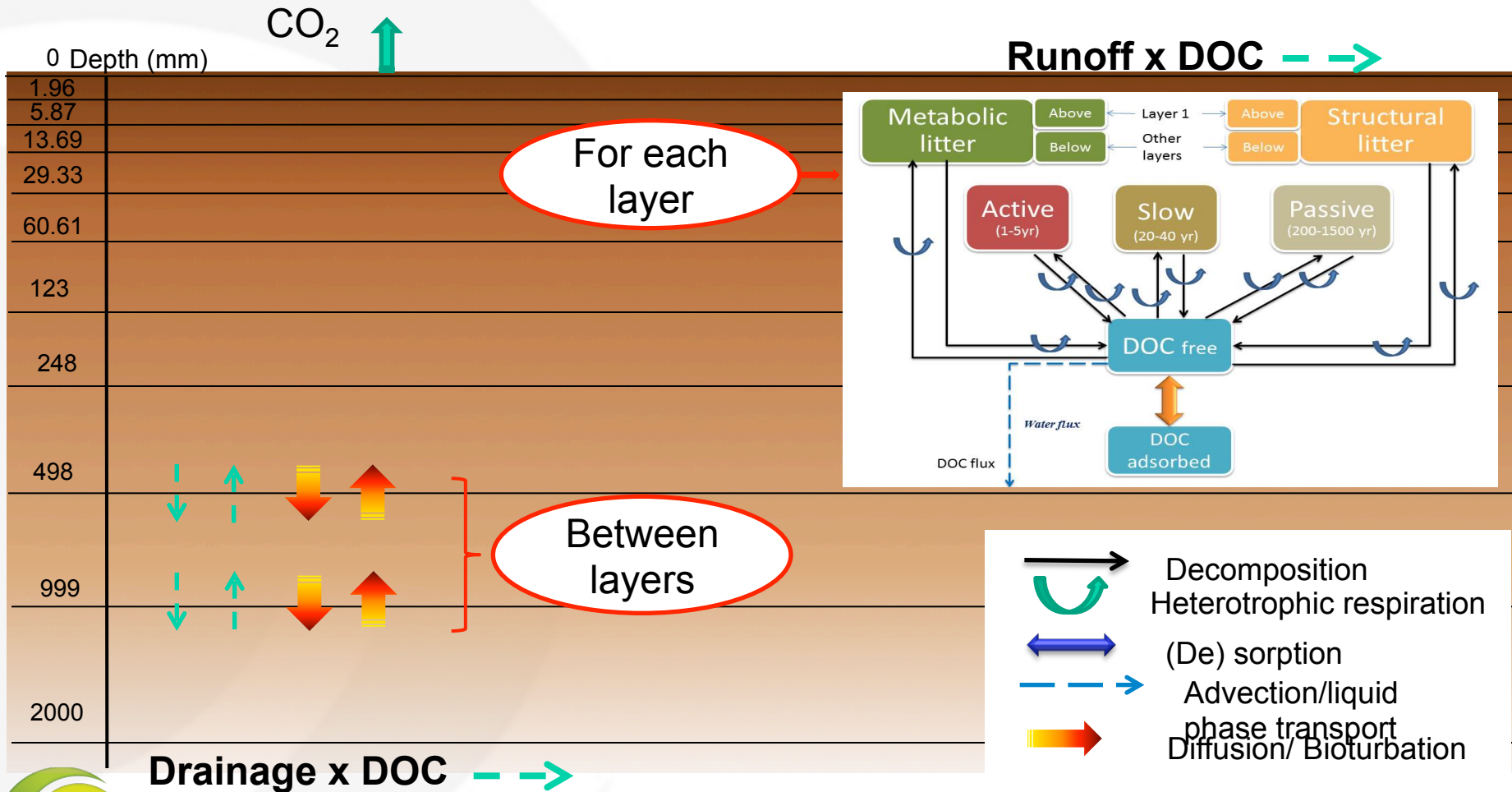
# Soil Carbon discretization

ORCHIDEE SVN r3340



# Soil Carbon discretization

Soil C discretized using the same layers than hydrology scheme (11 layers). A new pool introduced (DOC)



# Soil Carbon discretization

- Adsorption of DOC following initial mass isotherms

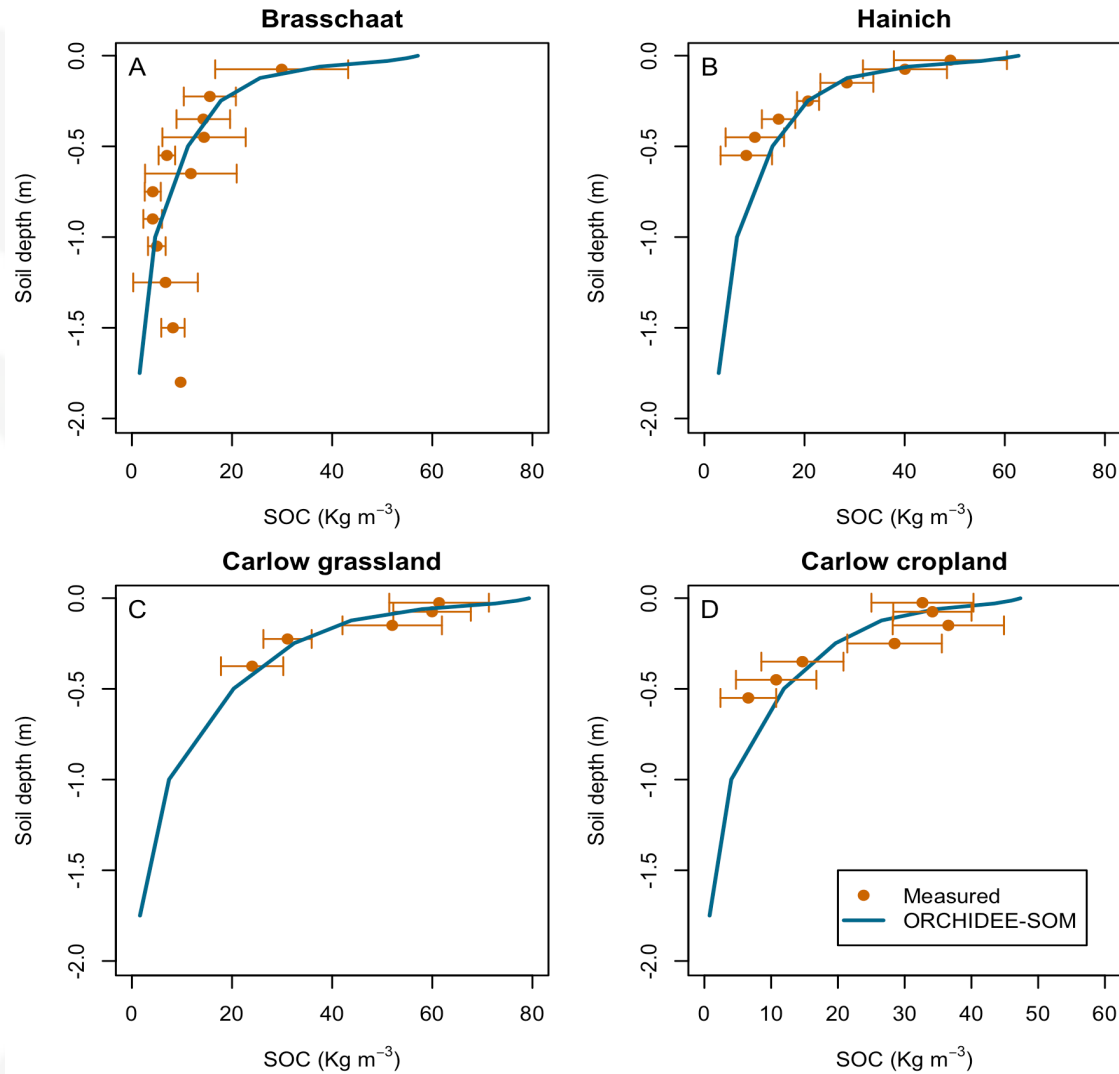
$$DOC_{ads} = Kd \times DOC_{free}$$

- DOC transported within the profile following the water movements (Futter et al., 2007) and exported following the runoff and the drainage fluxes
- POC and DOC transported using the second Fick's law

$$F_D = -D \times \frac{\partial^2 C}{\partial z^2}$$



# SITE SIMULATIONS



Camino-Serrano et al., (2018)

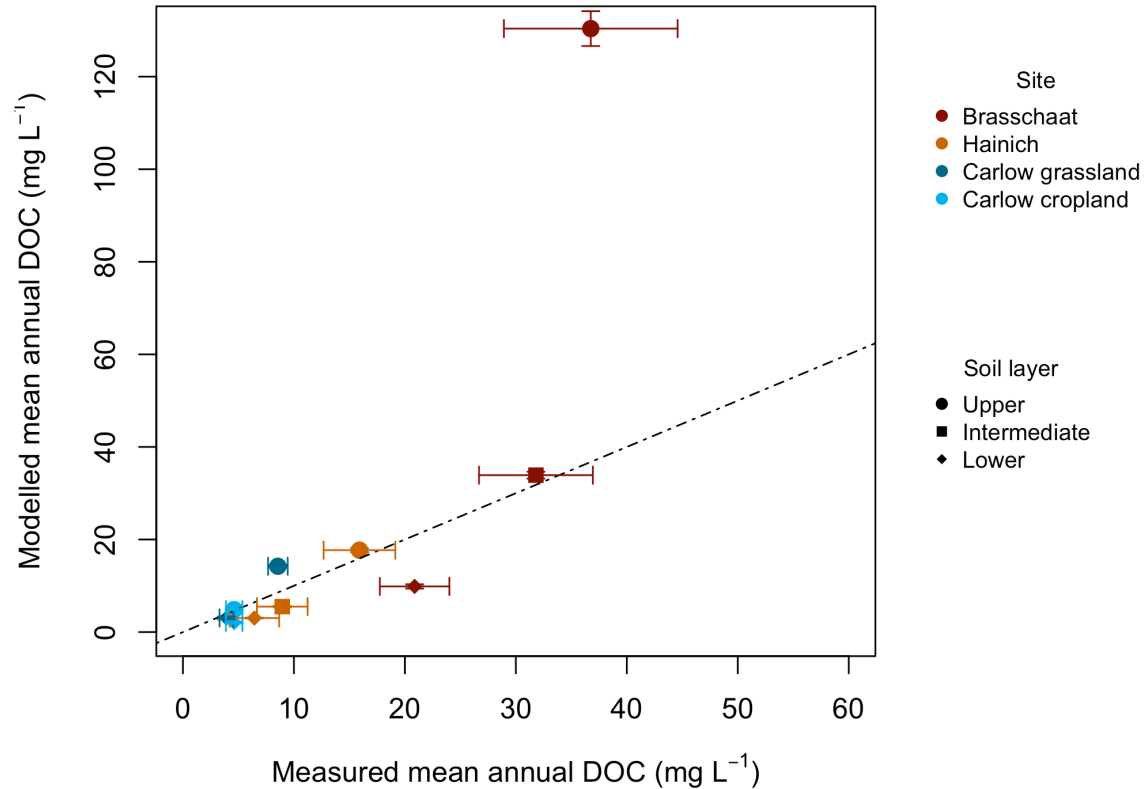


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# SITE SIMULATIONS

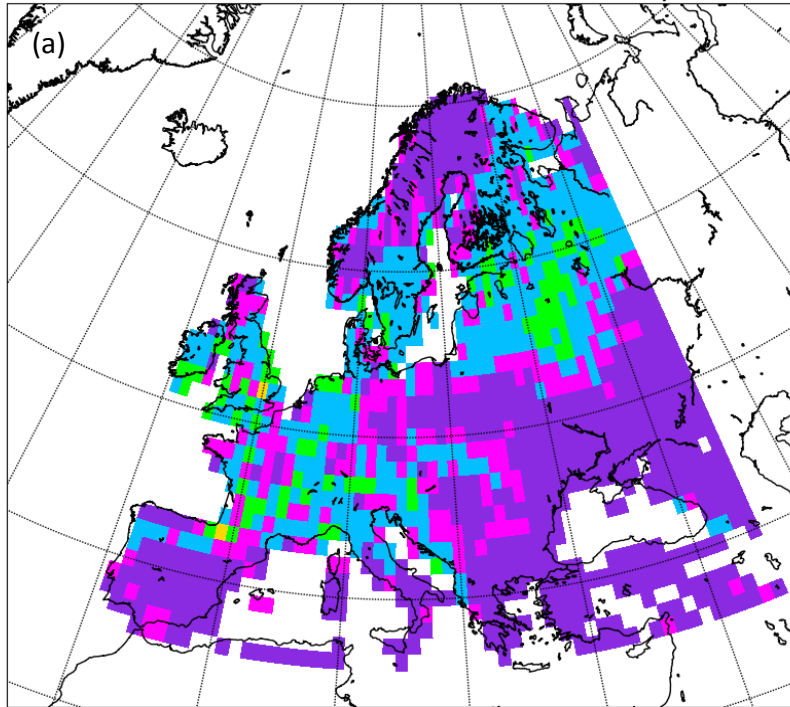


Camino-Serrano et al., (2018)

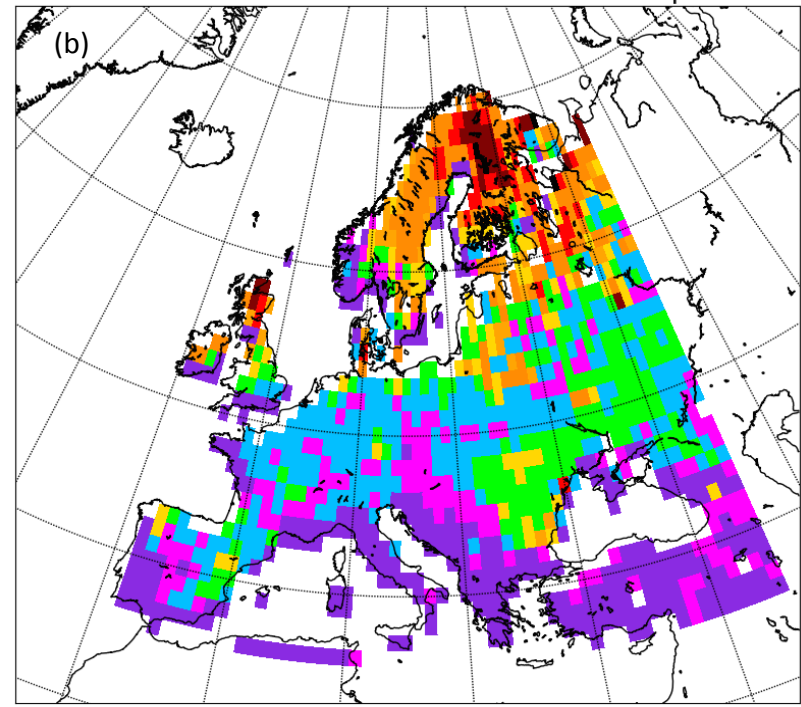


# OVER EUROPE

Total simulated SOC stock till 1.75m depth



Total SOC stock from GSDE till 1.75m depth



SOC stock (kg C m<sup>-2</sup>)

Guenet et al., (*In prep*)



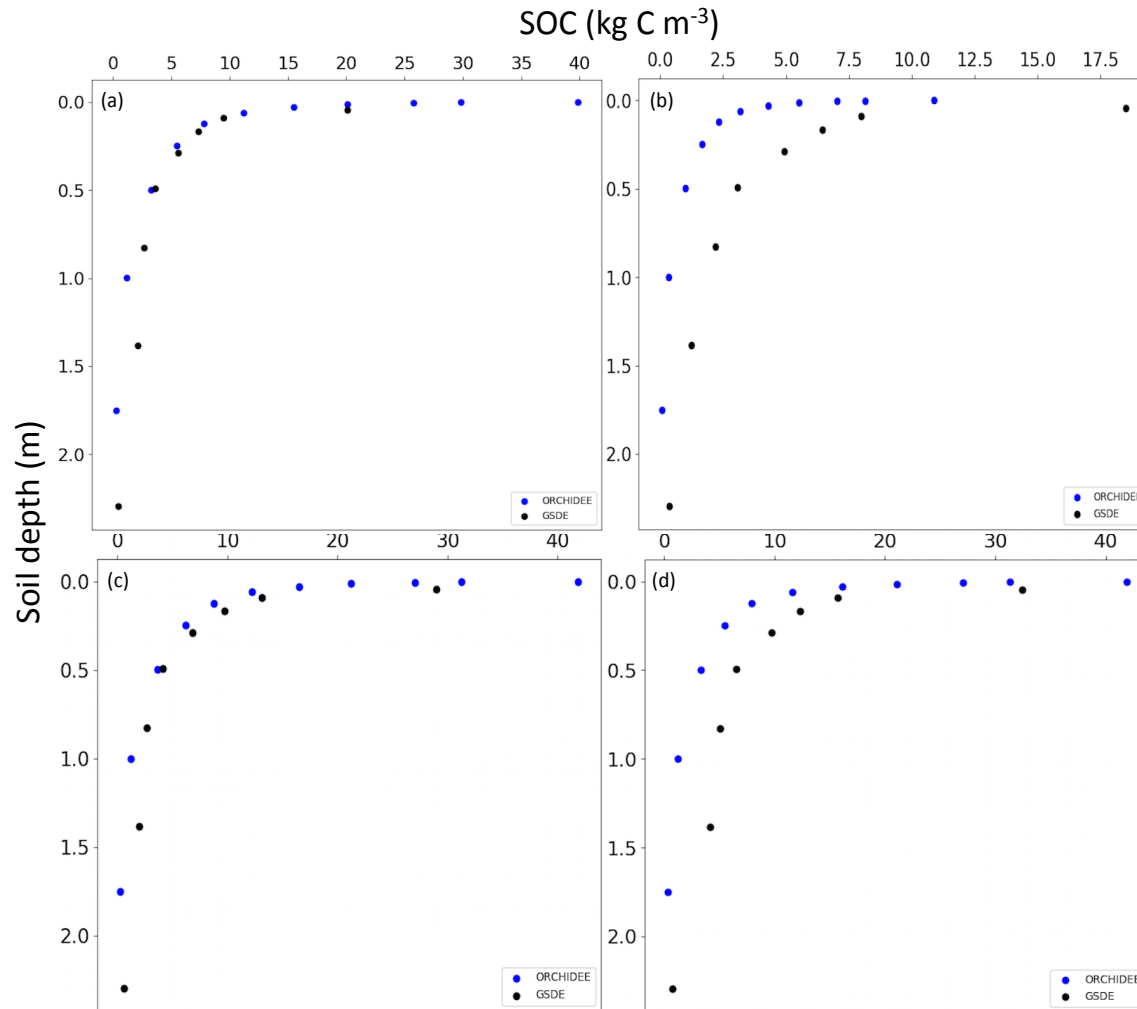
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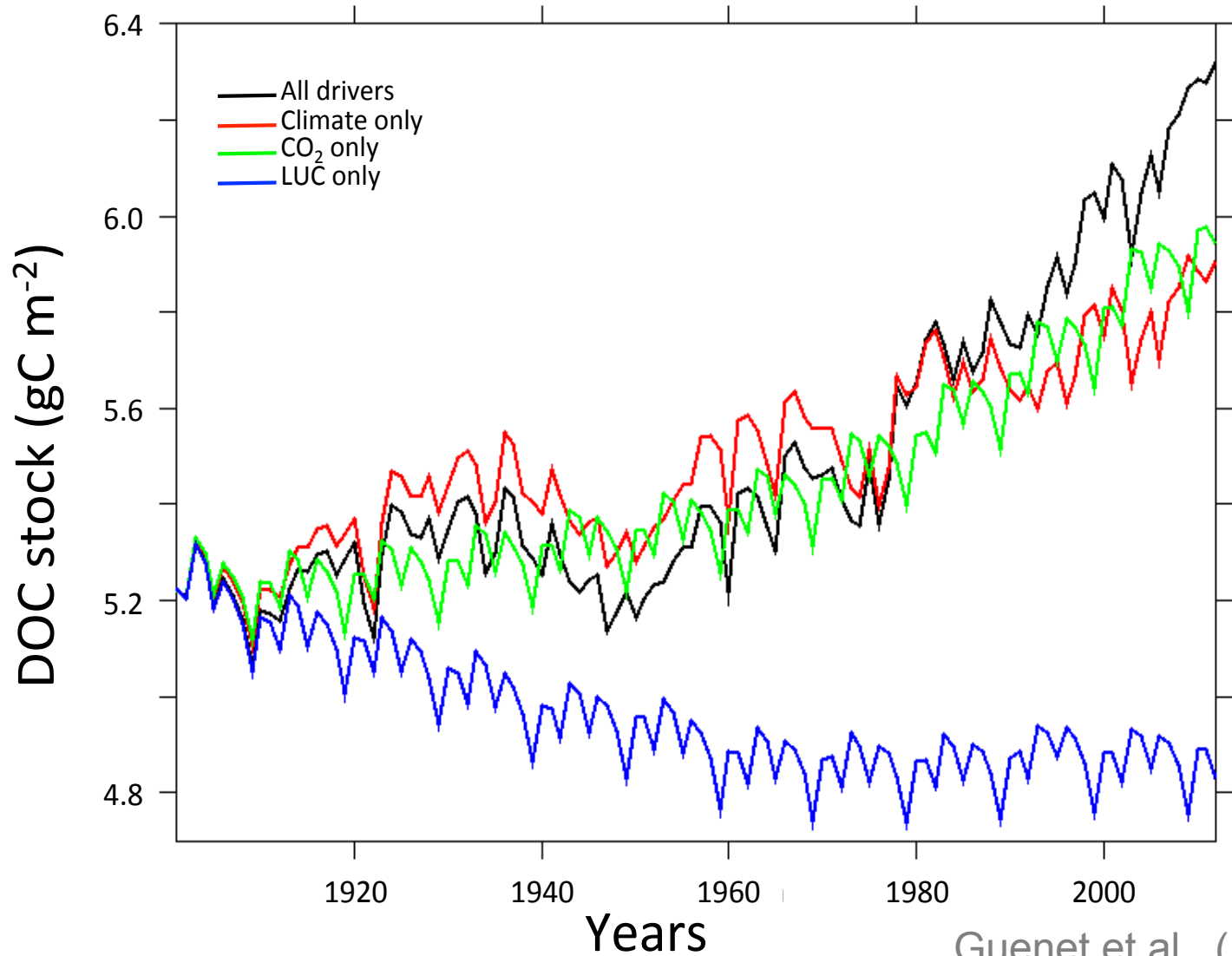
# OVER EUROPE



Guenet et al., (*In prep*)

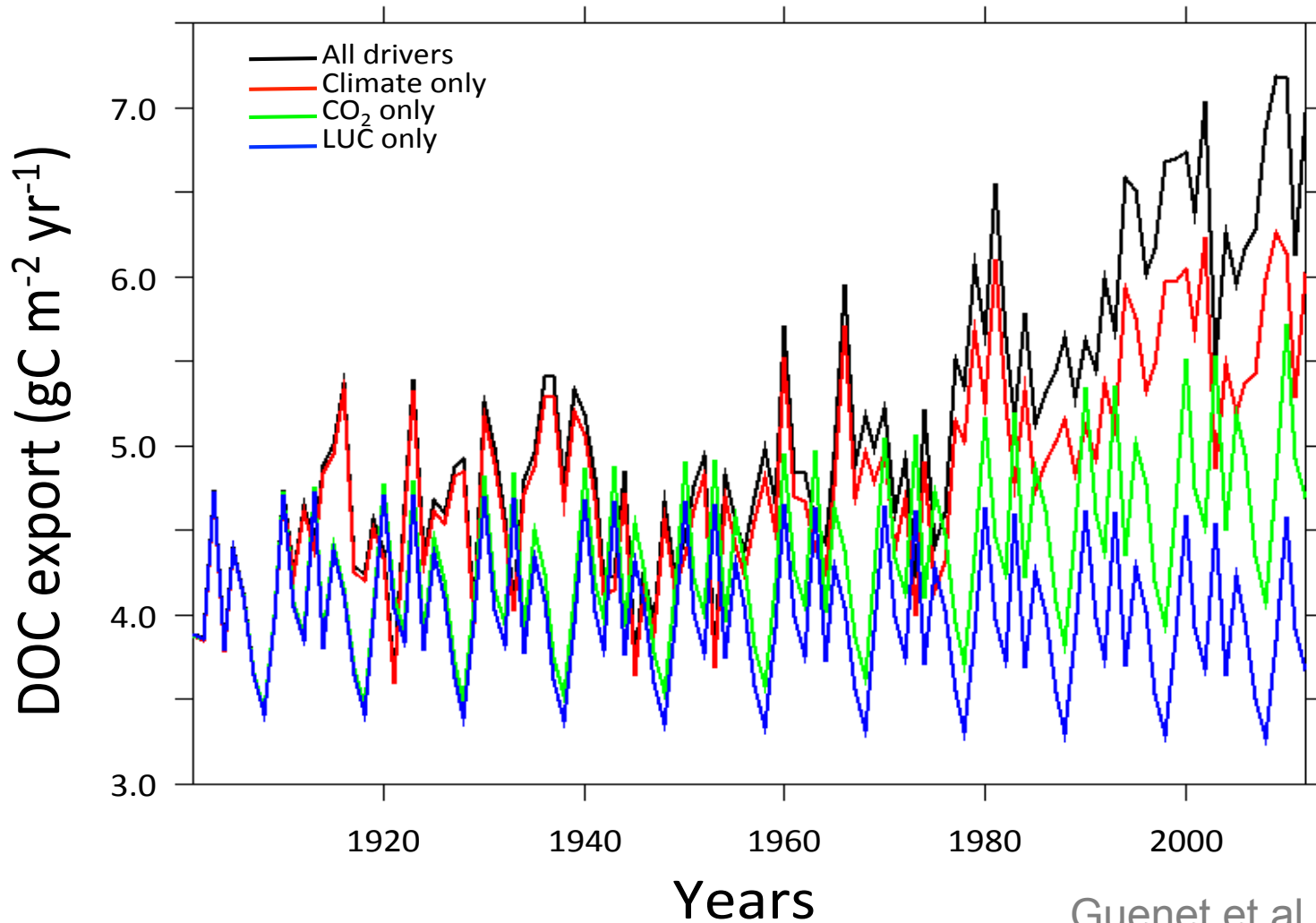


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Guenet et al., (In prep)

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# THANK YOU FOR YOUR ATTENTION!



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ORCHIDEE—training Jan. 2020

