

ORCHIDEE DEV: 18 novembre, 2014

The Soil Thermal Dynamics in ORCHIDEE

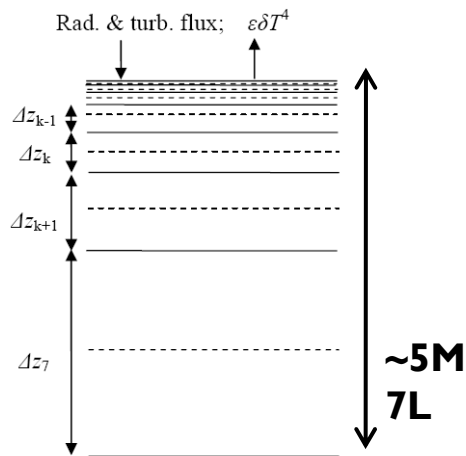
Fuxing WANG
LMD

Soil Thermal Model in ORC11

- ▶ **Current ORC-11 soil model:** (1) soil moisture changes with soil textures, but soil thermal properties do not; (2) different soil vertical discretization for moisture and temperature; (3) soil heat convection is neglected. **Energy imbalance !**

Soil Thermal [F. Hourdin, 1992]

$$C_p(\theta) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda(\theta) \frac{\partial T}{\partial z} \right]$$

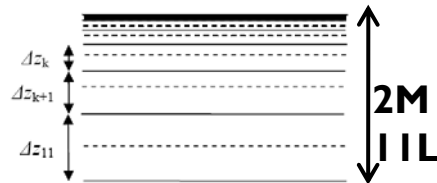


C_p : soil heat capa. (J/m³/K),
 λ : thermal conduc. (W/m/K).
 1 soil texture.

Soil Hydrology [P. de Rosnay, 1999]

$$\frac{\partial \theta}{\partial t} = \frac{\partial q_L}{\partial z} - s$$

$$q_L = -D(\theta) \frac{\partial \theta}{\partial z} + K(\theta)$$



θ : soil moisture (m³/m³);
 q_L : liquid water flux (m/s);

3 Soil Textures for Moisture



Revised

Revision (1):
 C_p and λ change with 3 soil textures & moisture!

Soil Thermal Properties

Soil Heat Capacity:

Currently: $C_p(\theta) = C_{dry} + wet(C_{wet} - C_{dry})$

$$wet = \frac{\theta - \theta_w}{\theta_f - \theta_w}$$

$$\kappa(\theta) = \kappa_{dry} + wet(\kappa_{wet} - \kappa_{dry})$$

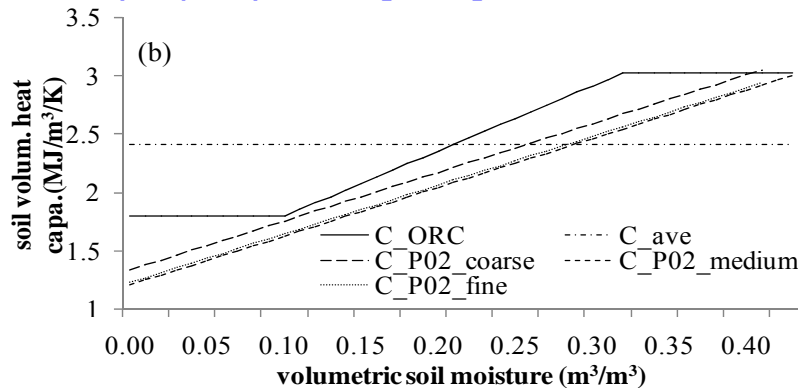
Soil Thermal Conductivity:

Soil thermal properties change with soil moisture, but independent of soil texture.

Revised:

$$C_p(\theta, st) = C_d(st) + \theta C_w$$

Dry capacity, Pielke [2002]; Volumetric SM



Soil thermal properties change with soil texture & soil moisture.

$$\kappa(\theta, st) = Ke \times (\kappa_{sat} - \kappa_{dry}) + \kappa_{dry}$$

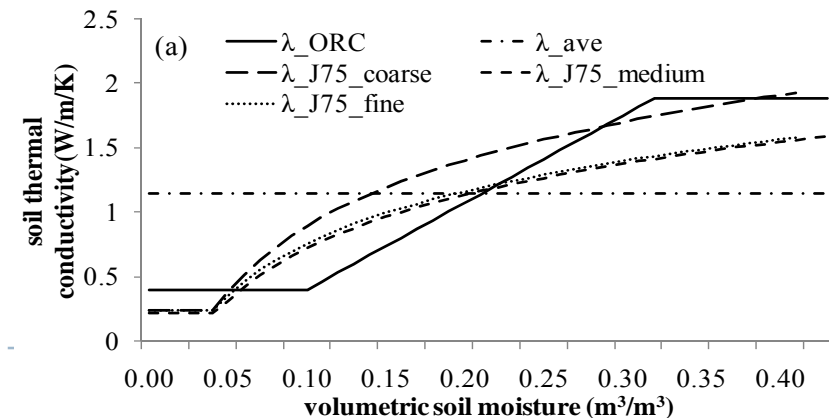
$$\kappa_{dry} = \frac{0.135 \gamma_d + 64.7}{2700 - 0.947 \gamma_d} \quad \gamma_d = (1 - n_p) \times 2700$$

$$\kappa_{sat} = \kappa_s^{1-n} \kappa_i^{n-x_u} \kappa_w^{x_u} \quad \kappa_s = \kappa_q^q \kappa_o^{1-q}$$

$$Ke = 0.7 \log Sr + 1.0$$

Saturation degree

Johansen, O. [1975, J75], Peters-Lidard et al. [1998]



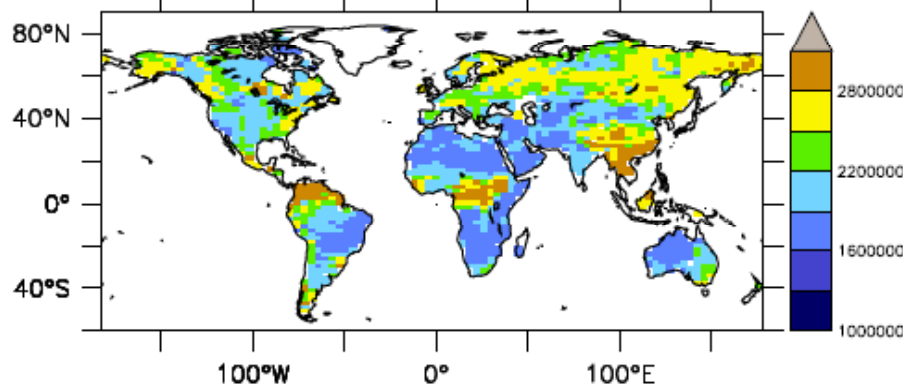
Soil Thermal Properties in ORC-LMDZ JJA (3-Year AVE)

Item	λ, C_p	Layer
CTL	original	original
EXP	revised	original

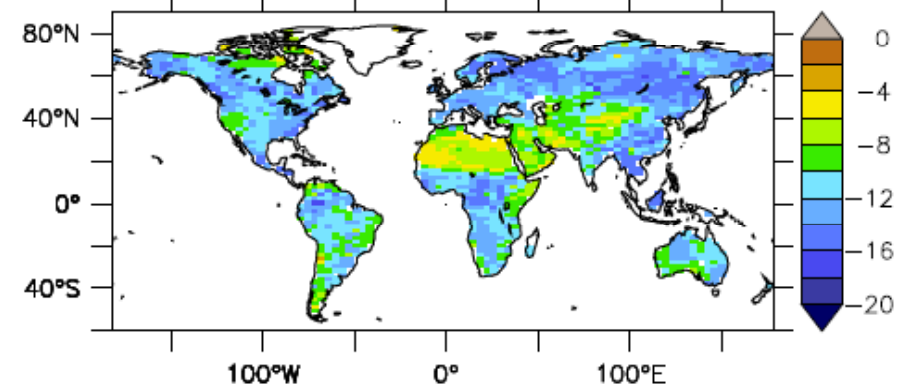
► **Standard**

Revised – Standard, %

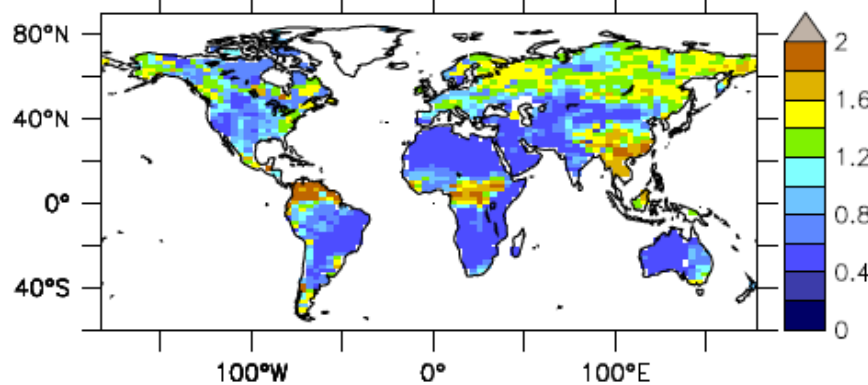
(c) SOILCAPA CTL(J/m³/K,JJA)



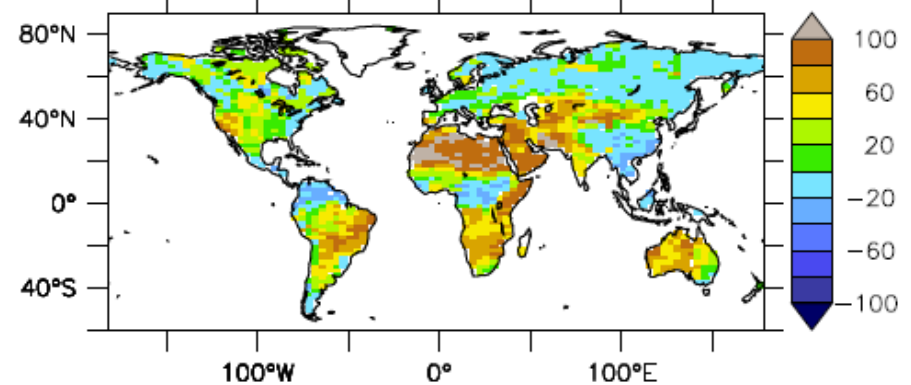
(d) SOILCAPA EXP-CTL(,JJA)



(e) SOILCOND CTL(W/m/K,JJA)



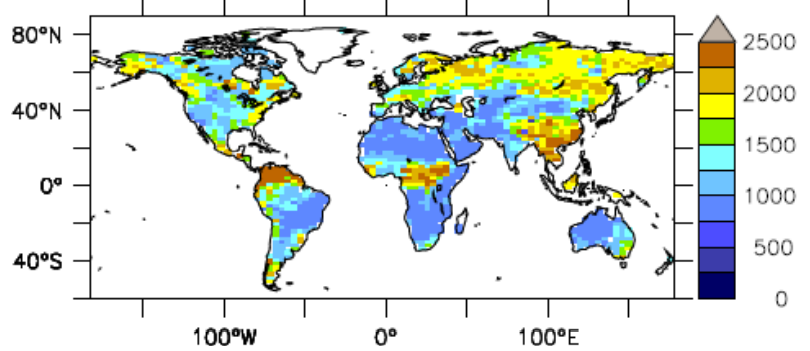
(f) SOILCOND EXP-CTL(,JJA)



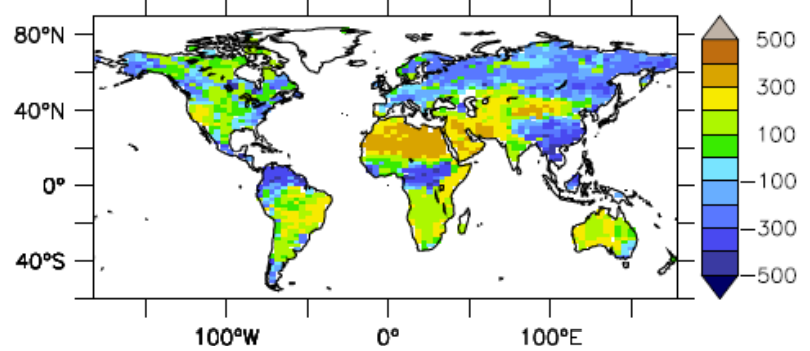
Effects of Soil Thermal Properties in ORC-LMDZ, JJA (3-Year AVE)

Definition: Soil thermal inertia, $P = (\lambda * C_p)^{0.5}$, ($W/m^2/K*s^{0.5}$), The resistance of soil to temp. change. The higher the P is, the slower the temperature varies during a full heating/cooling cycle (e.g., 24-hour day).

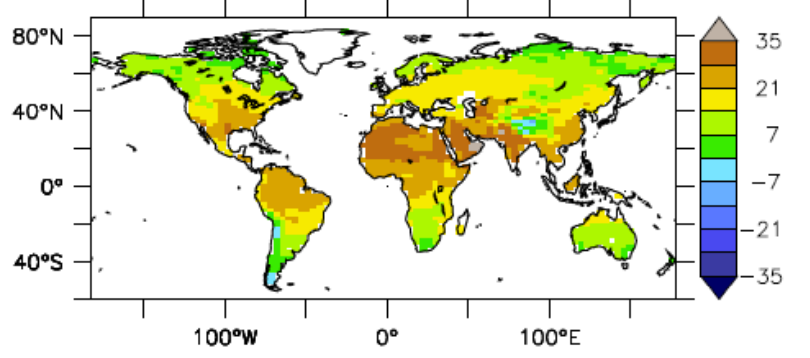
► **Standard**
(i) SOILINER CTL($W/m^2/K*s^{0.5}$,JJA)



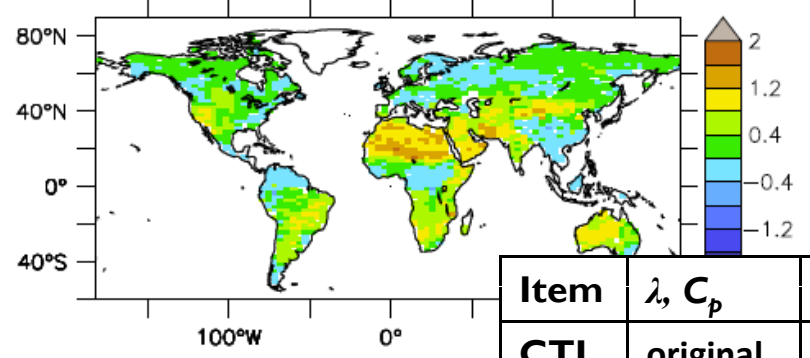
Revised - Standard
(j) SOILINER EXP-CTL($W/m^2/K*s^{0.5}$,JJA)



(k) TSURF CTL(Celsius,JJA)



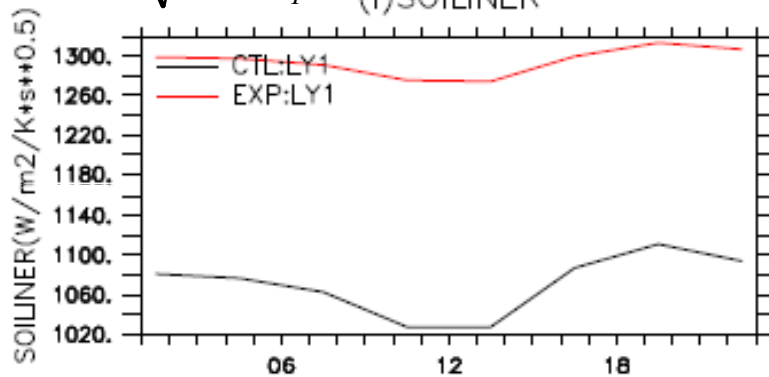
(l) TSURF EXP-CTL(Celsius,JJA)



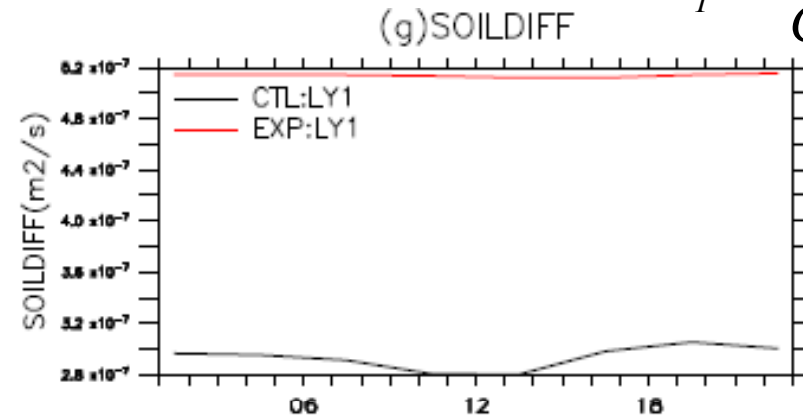
Item	λ, C_p	Layer
CTL	original	original
EXP	revised	original

Soil Thermal Prop. Effect: Diurnal Cycle

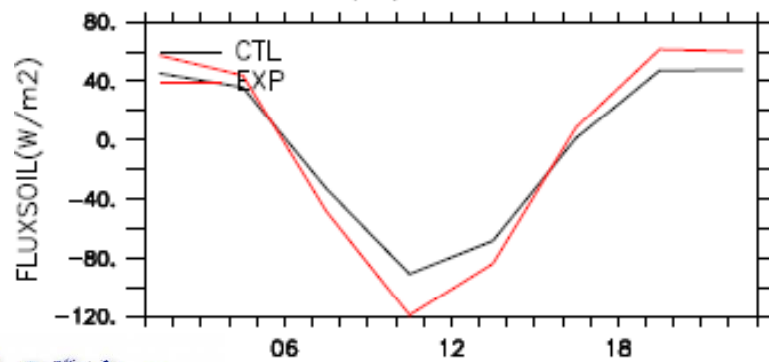
$$P = \sqrt{\lambda \cdot C_p} \quad \text{(i)SOILINER}$$



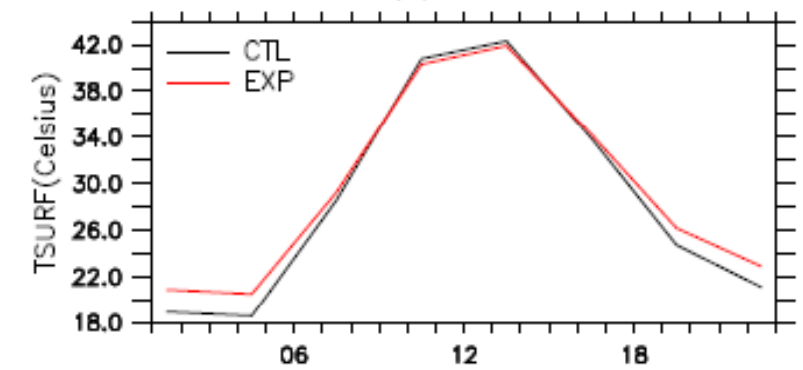
$$K_T = \frac{\lambda}{C_p}$$



(m)FLUXSOIL



(k)TSURF

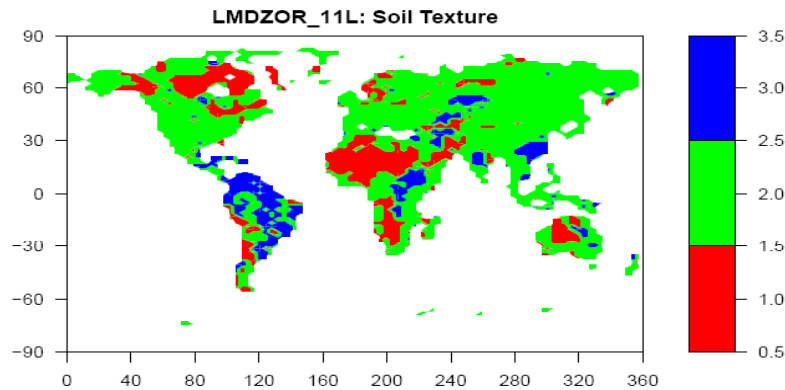


Sahara: 20W-5E, 10N-35N

Item	λ, C_p	Layer
CTL	original	original
EXP	revised	original

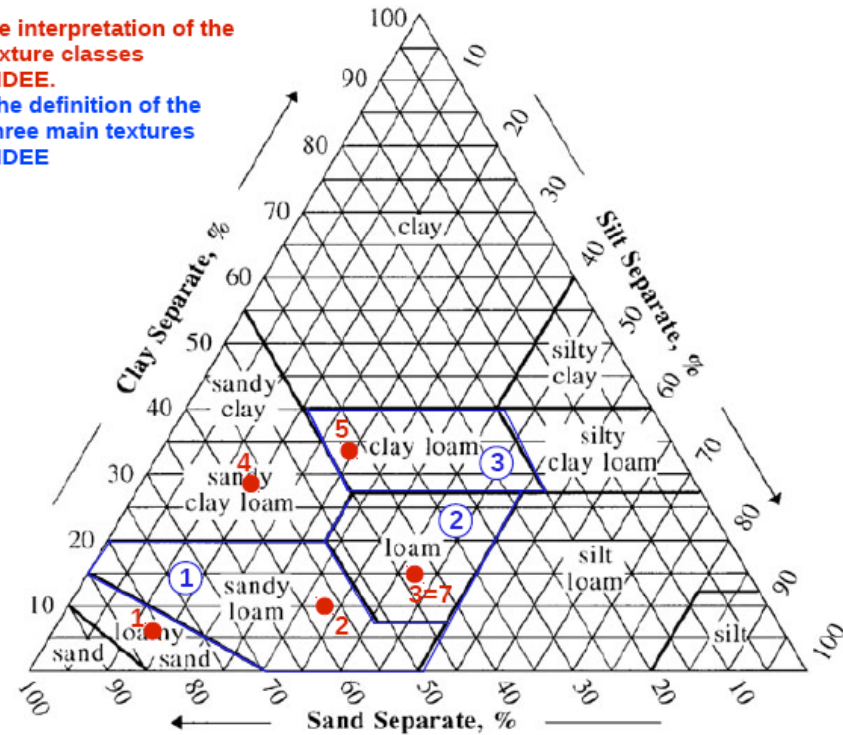
Soil Texture Classes

► Currently



Coarse (1), Medium (2), Fine (3)

In red, the interpretation of the Zobler texture classes in ORCHIDEE.
 In blue, the definition of the default three main textures in ORCHIDEE



USDA triangle and its 12 soil texture classes [Agnes Ducharne, Feb. 2014]

**USDA 12 Soil texture map in preparation;
 Soil thermal property parameterization for 12 classes done !**

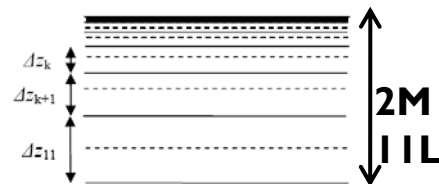
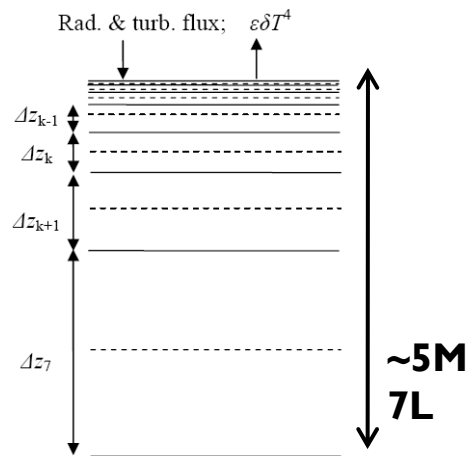
Soil Thermal Model in ORC11

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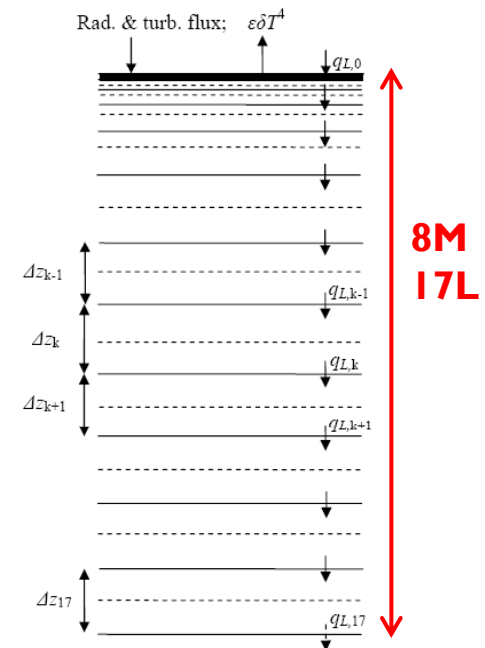
Soil Thermal [F. Hourdin, 1992]

Soil Hydrology [P. de Rosnay, 1999]

Revised



Revision (2)



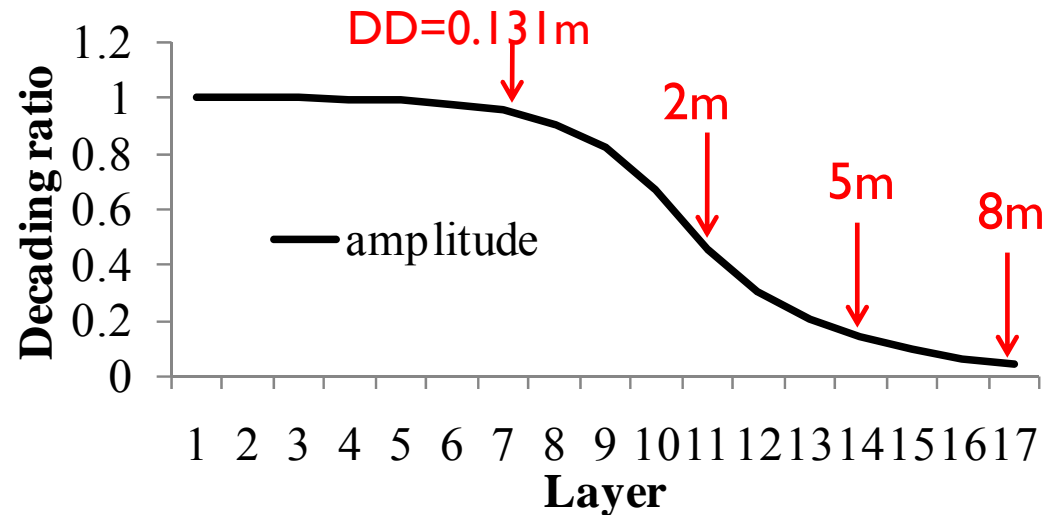
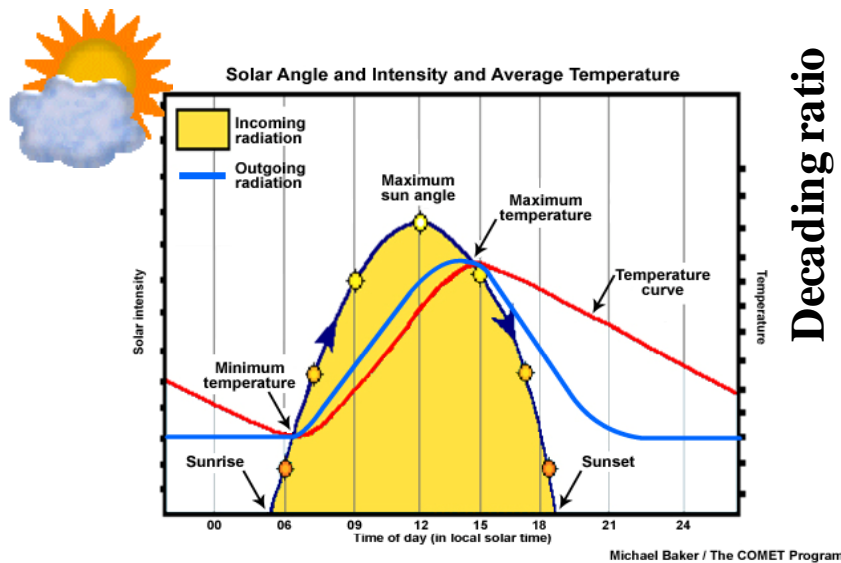
— Layer
 - - - Node

Soil Layer Thickness

layer	5M7L		8M17L	
	zz (m)	zl (m)	zz (m)	zl (m)
1	1.419E-2	3.426E-2	0	0.978E-3
2	6.264E-2	1.028E-1	1.955E-3	3.910E-3
3	1.595E-1	2.398E-1	5.865E-3	9.775E-3
4	3.533E-1	5.139E-1	1.369E-2	2.151E-2
5	7.409E-1	1.062	2.933E-2	4.497E-2
6	1.516	2.158	6.061E-2	9.189E-2
7	3.066	4.351	1.232E-1	1.857E-1
8			2.483E-1	3.734E-1
9			4.985E-1	7.488E-1
10			9.990E-1	1.500
11			2.000	<u>2.00/2.50</u>
12			<u>3.001</u>	<u>3.501</u>
13			<u>4.002</u>	<u>4.502</u>
14			<u>5.003</u>	<u>5.503</u>
15			<u>6.004</u>	<u>6.504</u>
16			<u>7.005</u>	<u>7.505</u>
17			<u>8.006</u>	<u>8.006</u>

Why Extend Soil Depth to 8M ?

- ▶ The amplitude of soil temperature decreases with depth.



$$\lambda = 1.329 \text{ W/m}^2/\text{K}, C = 2.135 \text{ J/m}^3/\text{K} \text{ (medium wetness)}$$

Damping Depth (DD): at which the amplitude decreases to the fraction $1/e$ of the surface amplitude. The required soil depth for

Diurnal Cycle: $3DD \approx 0.393\text{m}$

Annual Cycle: $3DD * \sqrt{365} \approx 7.500 \text{ m}$

Analytical Solution: Soil Heat Conduction

$$C_p \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda \frac{\partial T}{\partial z} \right]$$

- ▶ The **analytical solution** can be obtained by the **boundary/initial conditions**:

$$T(0, t) = \bar{T} + A_0 \sin(\omega t - \Phi)$$

$$\lim_{z \rightarrow \infty} T(z, t) = \bar{T}$$

$$T(z, 0) = \bar{T} + A_0 e^{-\frac{z}{D}} \sin\left(-\frac{z}{D} - \Phi\right)$$



$$T(z, t) = \bar{T} + A_0 e^{-\frac{z}{D}} \sin\left(\omega t - \frac{z}{D} - \Phi\right)$$

Angular velocity of earth rotation

$$\omega = \frac{2\pi}{\tau}$$

$$D = \sqrt{\frac{2K_T}{\omega}} = \sqrt{\frac{K_T \tau}{\pi}}$$

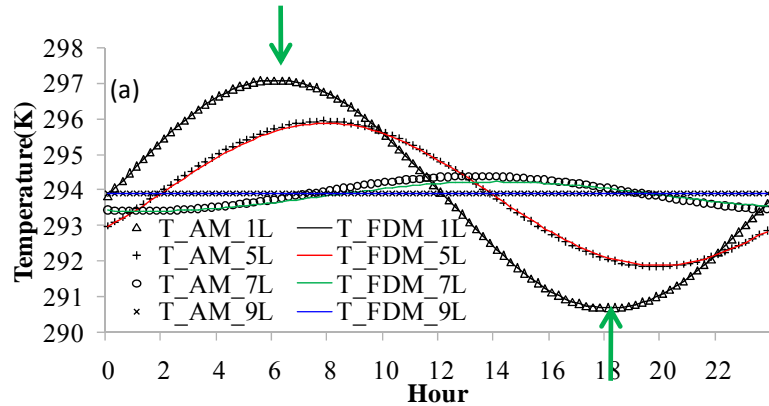
$$K_T = \frac{\lambda}{C_p}$$

Harmonic period of surface temperature (86400s diurnal cycle)

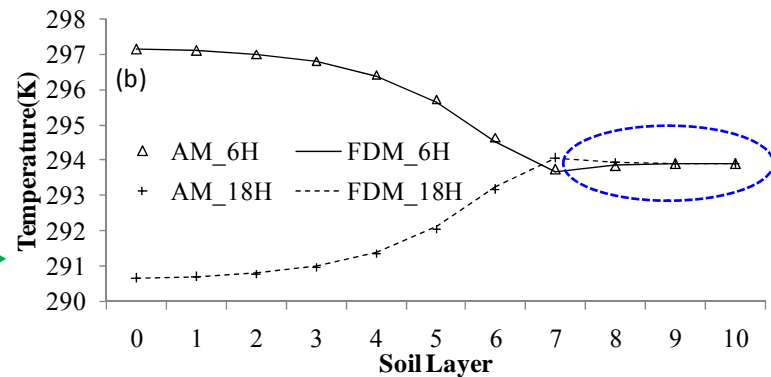
Numerical (FDM) vs. Analytical Solution (1D)

$$C_p \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda \frac{\partial T}{\partial z} \right]$$

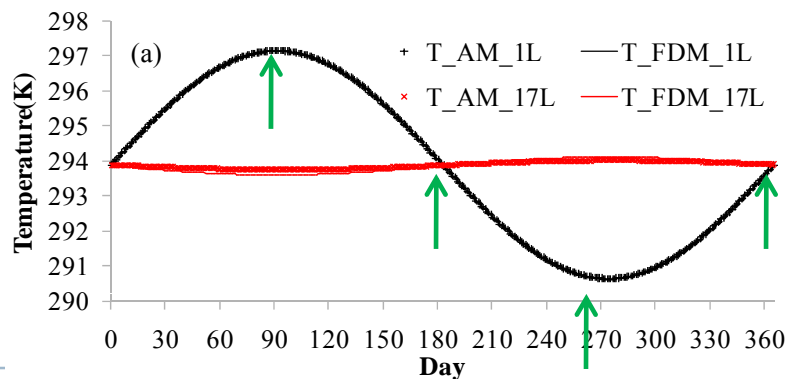
Diurnal Cycle: Time series at different layers



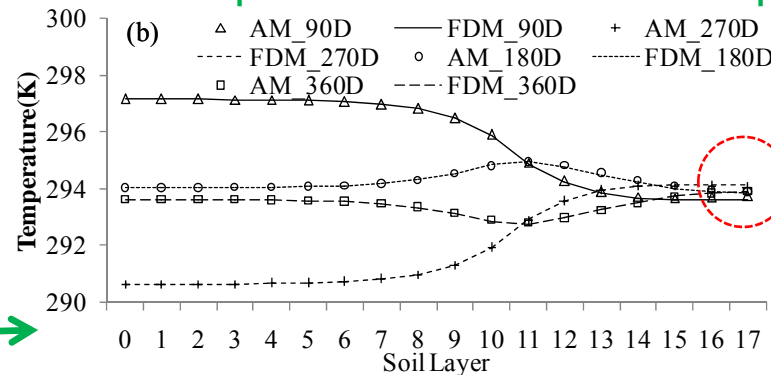
Vertical profile at different time steps



Annual Cycle: Time series at different layers

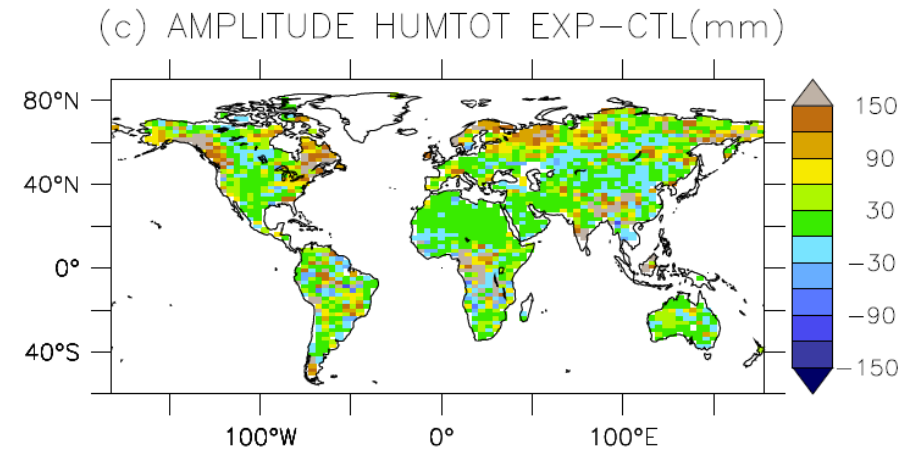
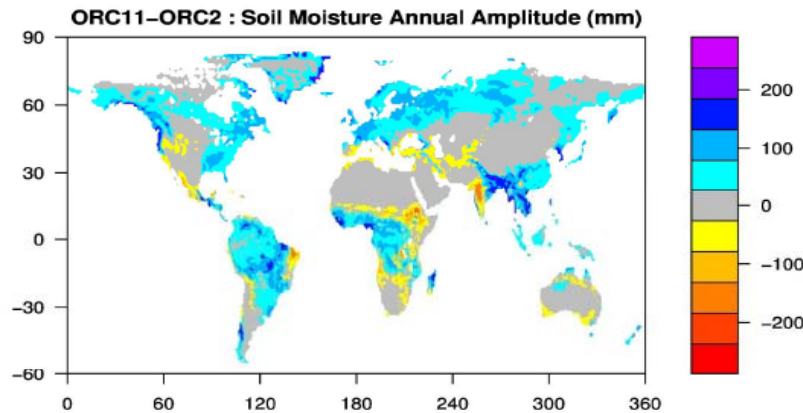


Vertical profile at different time steps

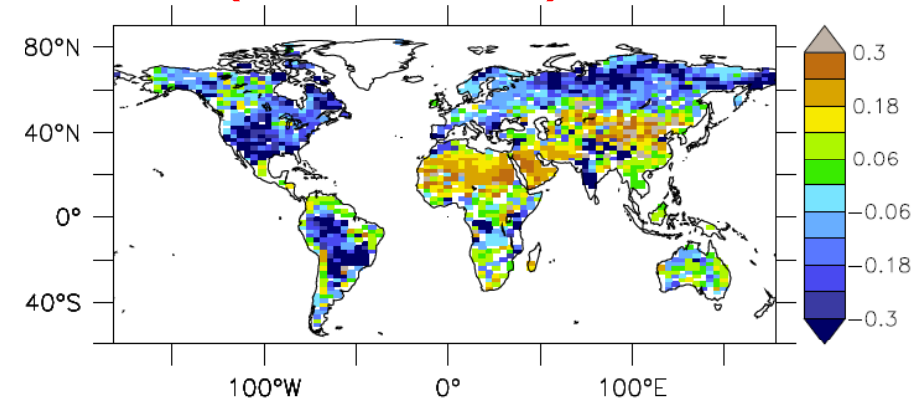
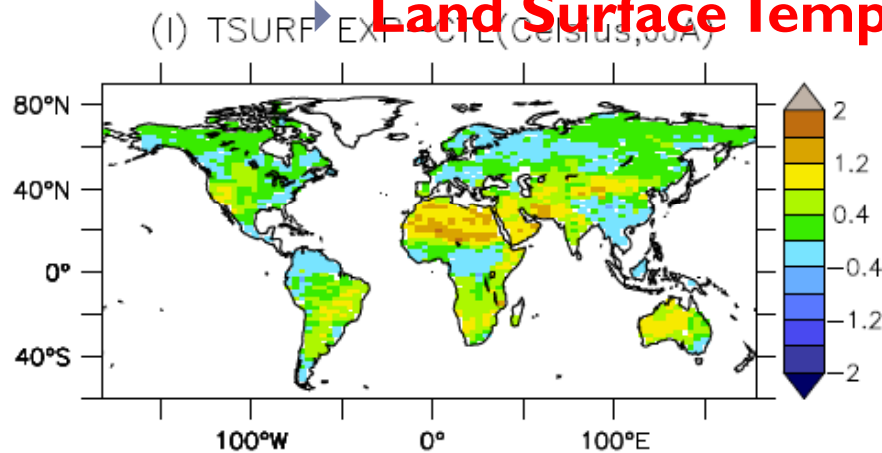


Discretization Effects (3-Year; SM,T)

► Moisture Annual Amplitude (mm)



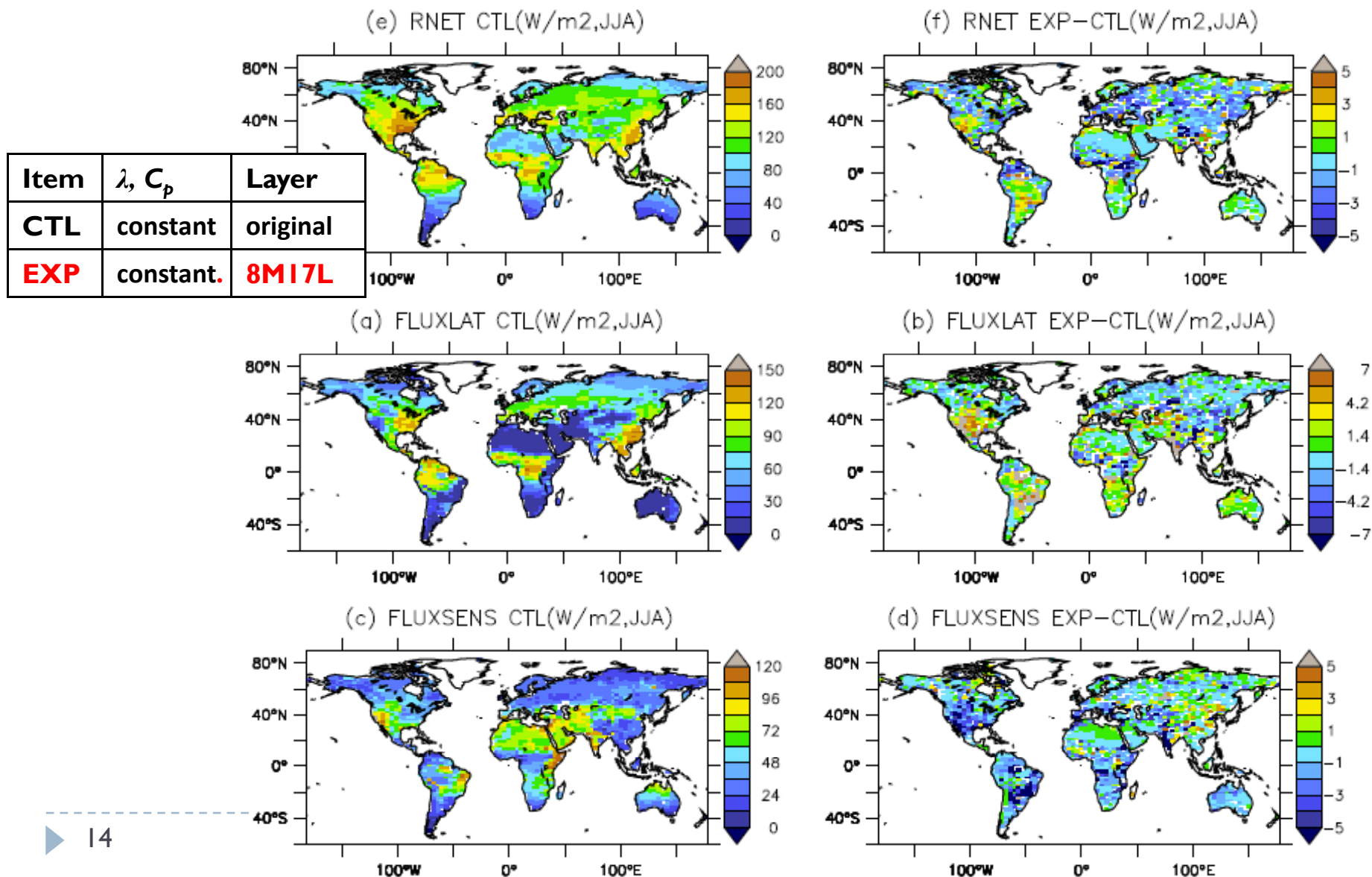
► Land Surface Temperature (Celsius, JJA)



► 13 Thermal property effects

Discretization effects

Discretization Effects (3-Year; Rn, LE, H)



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θ : soil moisture (m³/m³);
 q_L : liquid water flux (m/s);

Revised

Revision (3)

$$C_p(\theta, st) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda(\theta, st) \frac{\partial T}{\partial z} \right] + C_w \frac{\partial q_L T}{\partial z} - S$$

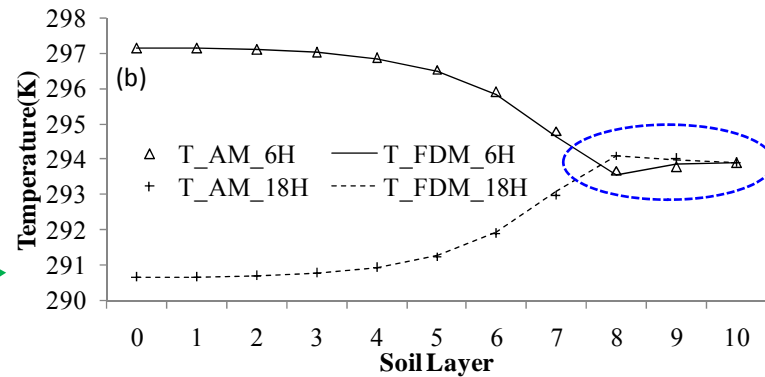
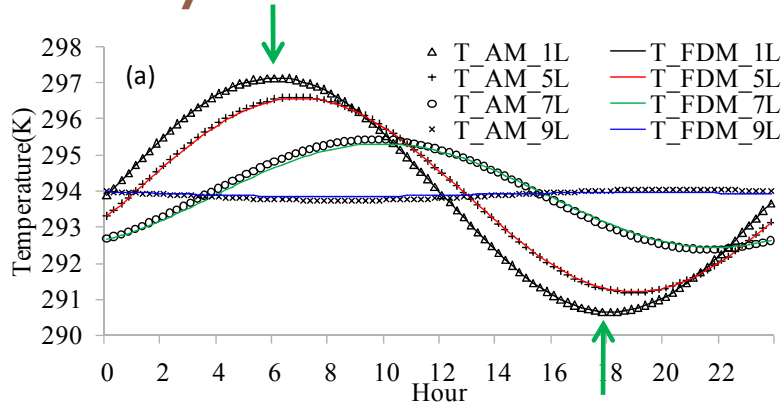


Numerical (FDM) vs. Analytical Solution Soil Heat Cond.-Convec. Steady Flow (1D)

$$C_p \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda \frac{\partial T}{\partial z} \right] - C_w \frac{\partial q_L T}{\partial z}$$

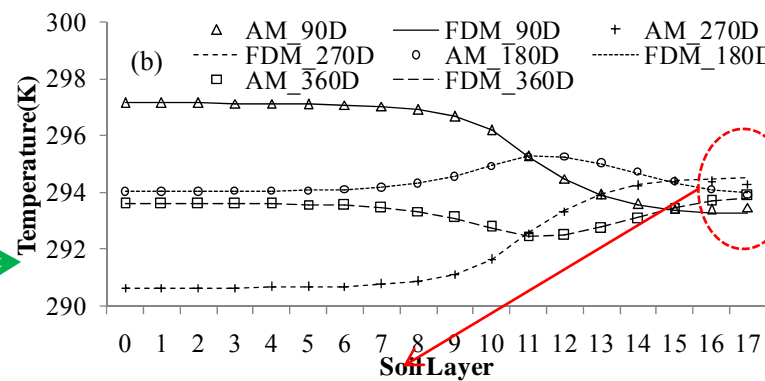
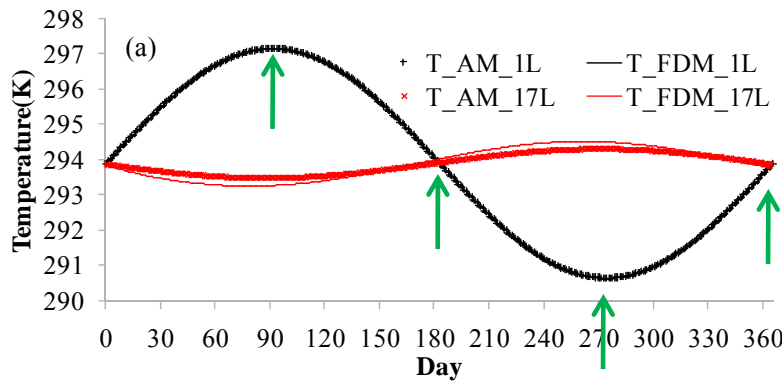
Diurnal Cycle: Time series at different layers

Vertical profile at different time steps



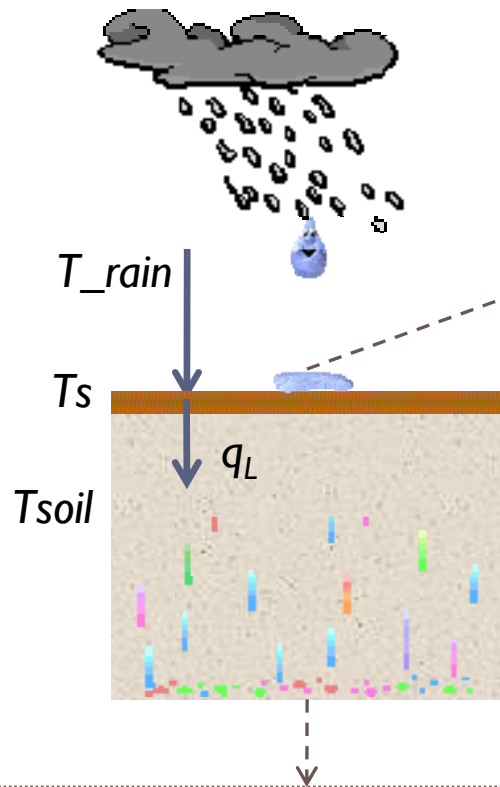
Annual Cycle: Time series at different layers

Vertical profile at different time steps



$q_L = 1E-7$ m/s (8.6 mm/d)

Soil Heat Conduction-Convection Coupling



Bottom BC: zero flux, $dT/dz=0$

Land Surface BC:

➤ Surface Energy Balance:

$$C_p \times dT_s/dt = R_{net} + LH + SH + G + \mathbf{H_{rain}}$$

$$\mathbf{H_{rain}} = C_w \times (T_{rain} - T_s) \times q_{L,0}$$

$$T_{rain} \approx T_{wetbulb}$$

$$T_{wetbulb} \sim f(T_{air}, Q_{air}, P_s)$$

$$C_p(\theta, st) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\underbrace{\lambda(\theta, st) \frac{\partial T}{\partial z}}_{\text{Conduction}} \right] - \underbrace{C_w \frac{\partial q_L T}{\partial z}}_{\text{Convection}} - S$$

The Effects of Liquid Water Transfer: ORCHIDEE Forced Mode, Daily Scale

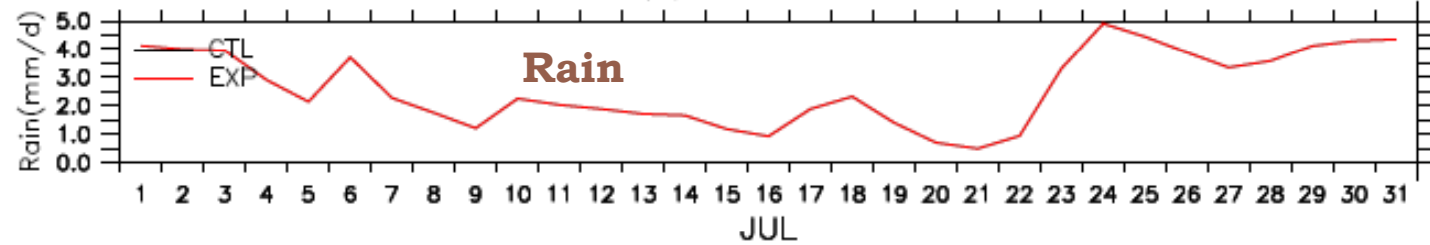


60W-110W,
25N-45N

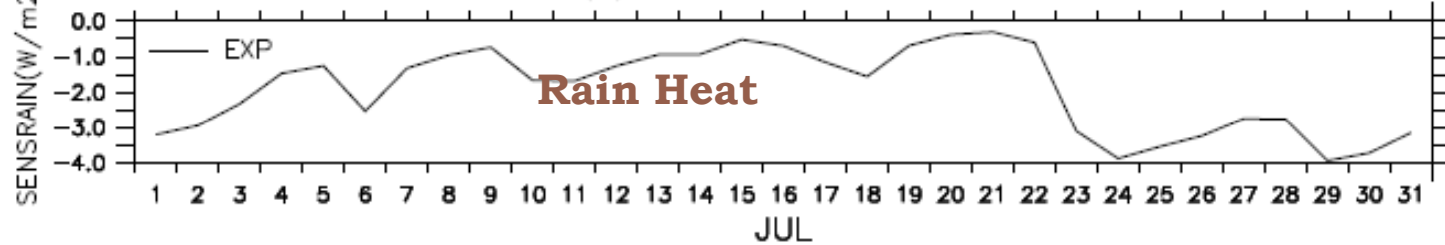
EXP-CTL: SURF
 Layer 3
 Layer 7
 Layer 10

EXP-CTL: FLUXSENS
 Positive upward

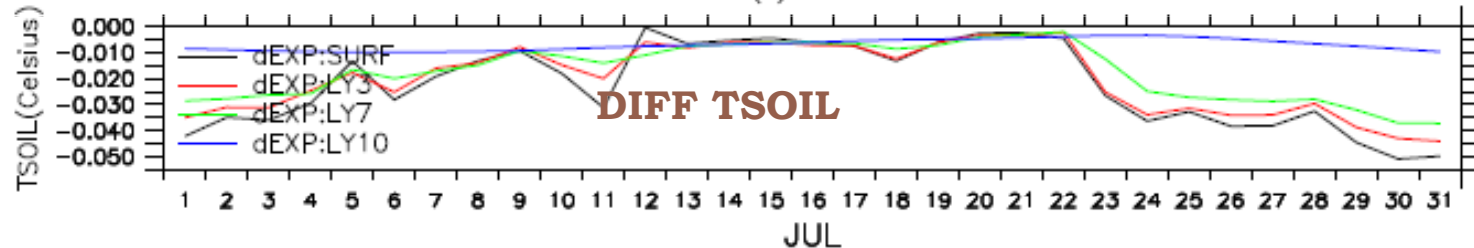
(b) Rain at Surf.



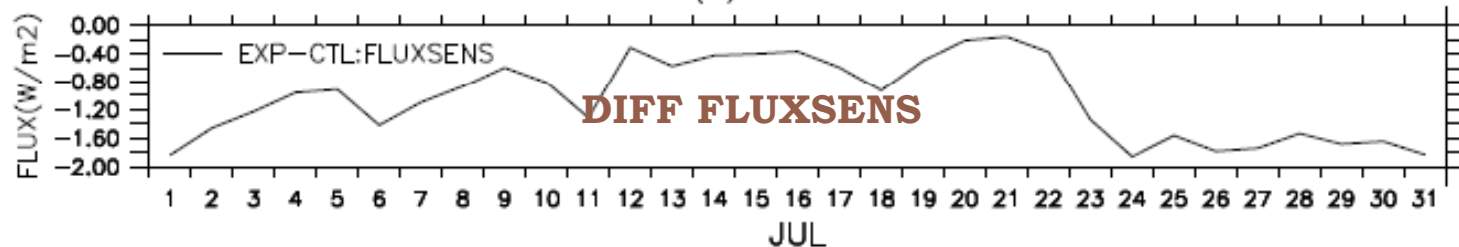
(d) Sensible Heat of Rain



(f) TSOIL

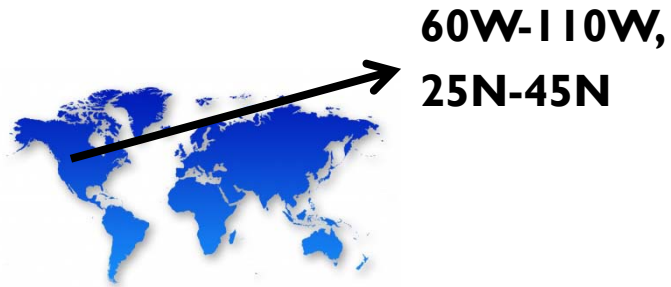


(n) FLUXSENS



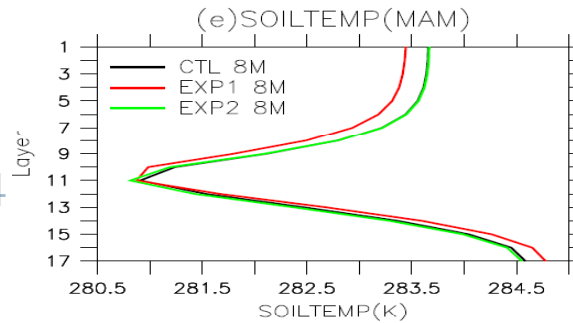
T_{soil} Vertical Profiles

Item	λ, C_p	q_L	Dep.
CTL 8M	Ave.	N	8M
EXP1 8M	Rev.	Y	17L
EXP2 8M	Ave.	Y	17L

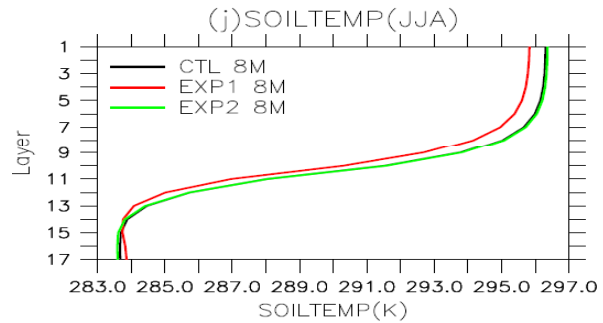


Conclusion: Soil thermal property (texture) effects are larger than q_L effects.

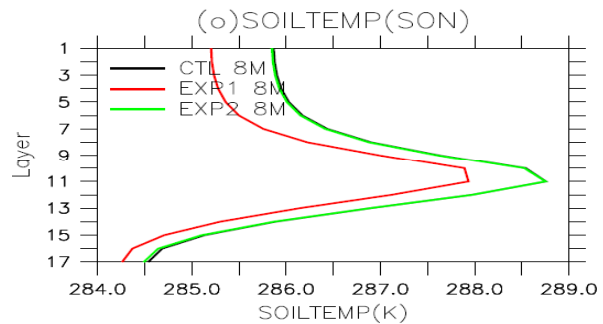
MAM



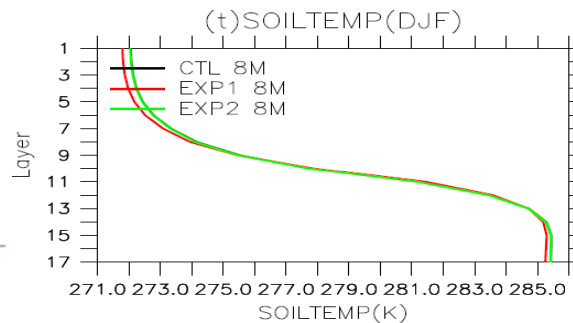
JJA



SON



DJF

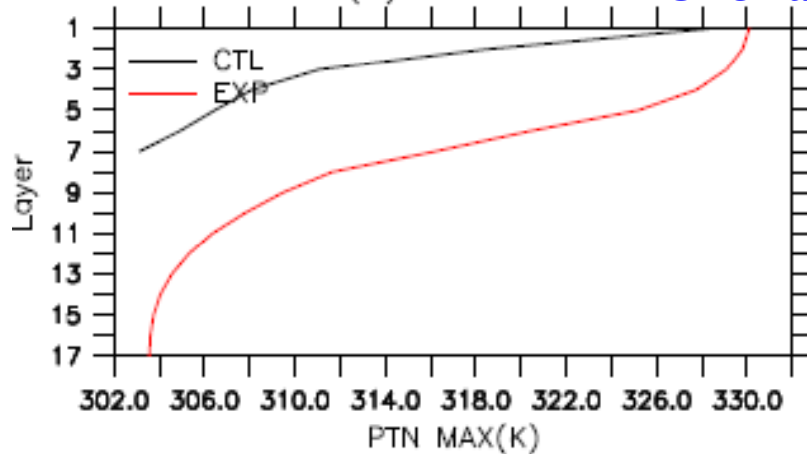


Check Numerical Stability in ORC-LMDZ (3-Hour Time Step, July)

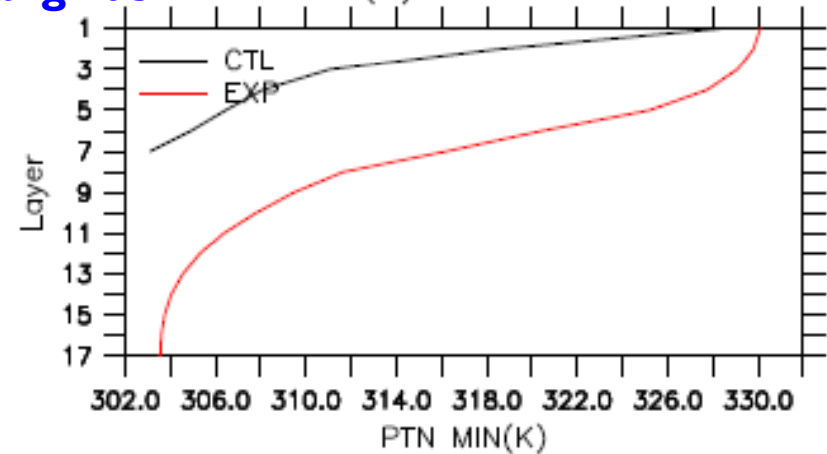
Item	λ, C_p	Liquid water (q_L)	Soil Depth
CTL	original	Not included	original
EXP	Revised.	Yes, included	8MI7L

Conclusion: (1) No extreme high/low temperature; (2) No oscillations in PTN, FLUXLAT.

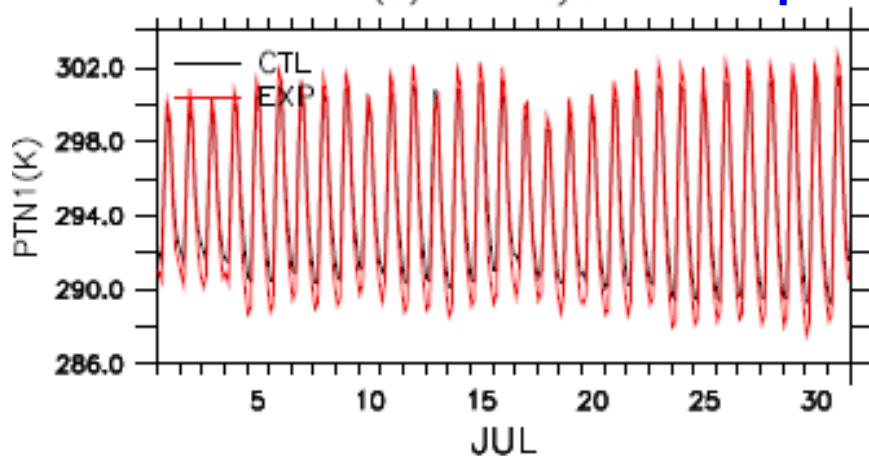
(a) PTN MAX **Over all global grids**



(b) PTN MIN



(c) PTN Layer1 **At point scale (25E,0N)**



(d) FLUXLAT

